

1 **Management of saltwater intrusion in coastal aquifers using different wells systems**
2 **considering climate change: Case study of the Nile delta**

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14 **ABSTRACT:**

15 Saltwater intrusion (SWI) is a special type of pollution that adversely affects the quality of
16 groundwater in coastal aquifers. The Nile Delta Aquifer (NDA) contains a large amount of
17 fresh water. Increasing abstraction from the aquifer and sea level rise have led to increased
18 SWI which reached to 100 km. Therefore, practical measures are required to prevent further
19 SWI. This study aims to identify an optimal well system to manage the intrusion of saline
20 water in NDA using a number of management systems including pumping of brackish water,
21 recharge of aquifer and abstraction of fresh water. The SEAWAT code is used to simulate SWI
22 in the aquifer considering different scenarios of pumping and sea level rise. Four scenarios are
23 used to control SWI including; decreasing pumping from the aquifer, increasing recharge
24 using treated waste water, increasing abstraction of brackish water for desalination and
25 combination of these systems. The results show that increasing recharge could lead to greater
26 retardation of SWI (19.50%) than decreasing pumping (6.20%) and abstraction of brackish
27 water (5.90%). However, a combined well system of pumping, recharge and abstraction is
28 shown to be a more effective tool to control SWI in coastal aquifers with retardation
29 percentage of 21.30%.

30 **Keywords:** saltwater intrusion, control, pumping, recharge, abstraction, Nile Delta aquifer.

31 **1. Introduction**

32

33 The high rate of population growth over the past 100 years has increased the water demand in
34 different areas of the world. The increasing water demand has increased pumping from
35 groundwater aquifers which has in turn caused saltwater intrusion (SWI) in coastal aquifers.
36 In normal circumstances, the freshwater flows in the coast direction but over-pumping may
37 lead to reversal of the flow inland causing SWI (Bear et al., 1999). Also, climate change has
38 been increasing at alarming rates in the last few decades which has led to rise in sea levels.
39 The rise in mean sea level may occur due to thermal expansion of seas or melting of glaciers
40 and ice caps. According to the Intergovernmental Panel on Climate Change (IPCC, 1996),
41 over the 20th century, sea levels have risen about 1-2 cm/yr. The expected rise in sea level by
42 the end of the 21st century is between 2-8.8 cm/yr. (IPCC, 2001).

43

44 The main causes of saltwater intrusion include over-abstraction, changes in groundwater flow,
45 tidal effects, seismic waves, dispersion and sea level rise. One of the main impacts of the sea
46 level rise on coastal regions is increasing SWI. It also has a direct impact on groundwater
47 resources, soil salinity, and agricultural productivity in coastal zones. Saltwater intrusion into
48 coastal aquifers should be controlled in order to protect groundwater resources from
49 deterioration. The control of saltwater intrusion requires keeping an appropriate balance
50 between water abstracted from and recharged to the aquifer (Bear et al., 1999). Different
51 measures to control SWI were presented by Todd (1974) including: decreasing pumping rate,
52 relocation of wells, use of subsurface barriers, natural and artificial recharge, and abstraction
53 of saline water and disposal to the sea.

54

55 A number of numerical models have been used to improve understanding of the relevant
56 processes that cause SWI and find appropriate control methods. A number of studies have
57 investigated saltwater intrusion. However, a relatively small amount of research has been
58 carried out to study the control of saltwater intrusion (Abd-Elhamid and Javadi, 2008). In this
59 study, a number of effective methods to control SWI are examined and compared including;
60 reducing pumping rates, artificial recharge, abstraction of saline water and combination of
61 these methods. A comprehensive review of the developed models to study and control SWI is
62 presented below.

63

64 Reducing abstraction rates aims to obtain a sustainable yield while using other sources of
65 water to meet the demand for water. The control of SWI has been studied using number of
66 models by managing the abstraction (e.g. Scholze et al. (2002), Zhou et al. (2003) and
67 Bhattacharjya and Datta (2005)). Qahman and Larabi (2004) used the SEAWAT code to study
68 the problem of extensive SWI in the Gaza aquifer, in Palestine. Saltwater intrusion was
69 investigated under different scenarios of abstraction over the time. The study showed that the
70 aquifer is very sensitive to pumping which should be managed to protect the aquifer from
71 SWI. Narayan et al. (2006) used various pumping and recharge schemes to determine the
72 current and possible future extent of saltwater intrusion in the Burdekin aquifer in Australia
73 using the SUTRA code.

74

75 Artificial recharge helps to raise groundwater levels in aquifers. Surface spread of water from
76 rivers or canals, treated wastewater, and desalinated water are possible sources of water for
77 recharging unconfined aquifers, while recharge wells can be used for confined aquifers. The
78 management of SWI in coastal aquifers using artificial recharge was studied by a number of
79 researchers such as Narayan et al. (2006), Papadopoulou et al. (2005) and Kashef (1976).

80 Mahesha (1996a) simulated the effect of recharge wells on SWI in confined aquifers. Several
81 conditions were studied by changing well spacing and recharge duration. The study showed
82 that spacing of the wells and rate and duration of recharge can decrease SWI by repulsion of
83 the saline wedge.

84

85 The abstraction saline water and its disposal to the sea helps to decrease the mass of saline
86 water in coastal aquifers. This method was applied to control saltwater intrusion in coastal
87 aquifers by a number of researchers such as Maimone and Fitzgerald (2001) and Kacimov et
88 al. (2009). Sherif and Hamza (2001) studied the balance between saltwater and freshwater
89 using abstraction of brackish water from the transition zone using a 2-D finite element model.
90 Sherif and Kacimov (2008) examined different pumping scenarios and suggested that
91 pumping brackish water can reduce the intrusion of saline water. The main problem in
92 abstraction of brackish water is that the disposal of the saline water to the sea could affect the
93 marine life, fishing and tourism in these areas (Abd-Elhamid and Javadi, 2008). Johnson and
94 Sperry (2001) studied a number of methods to control SWI in different states of the United
95 States of America. In California, the saline water was abstracted and desalinated using
96 Reverse Osmosis. The treated water was mixed with groundwater to produce an appropriate
97 water quality for domestic use. In Los Angeles, treated waste water was used to recharge the
98 aquifer using a system of wells to protect coastal aquifers from SWI.

99

100 Combination of two or more of the above methods could give a better control of SWI.
101 Different combined methodologies to control SWI in coastal zones were used by a number of
102 researchers (e.g. Zhou et al. (2003); Hong et al. (2004); Narayan et al. (2006), Paniconi et al.
103 (2001); and Barrocu et al. (2004)). Mahesha (1996c) used a series of wells for abstracting
104 saline water in combination with freshwater recharge wells. The study indicated that the

105 combination of abstraction and recharge gives better results compared with the abstraction or
106 recharge method, leading to larger spaces between wells and smaller recharge rates. Abd-
107 Elhamid and Javadi (2011) introduced the ADR (Abstraction, Desalination and Recharge)
108 method to control SWI. The ADR method includes abstraction of saline water, desalination
109 and recharge to the aquifer. A finite element model was developed and used to simulate SWI.
110 The results indicated that the ADR method is more effective than the abstraction or recharge
111 method. A simulation-optimization model was developed by Javadi et al., (2012) to control
112 SWI using the combination of the abstraction and recharge techniques. The developed model
113 was applied to a confined coastal aquifer. The results showed that using the combined
114 recharge and abstraction system is significantly more effective in controlling SWI than using
115 abstraction or recharge wells separately. The combination of abstraction and recharge wells
116 gave lower cost and lower salt concentration in the aquifer.

117
118 Javadi et al., (2013) used the SUTRA code integrated with a Genetic Algorithm optimization
119 tool to identify the efficacy of different scenarios of controlling SWI in an unconfined costal
120 aquifer. The method was based on a combination of abstraction of saline water and recharge
121 of aquifer using surface ponds. The results confirmed that the combined system is more
122 effective in controlling SWI than using separate methods. Javadi et al., (2015) presented a
123 methodology based on the combination of abstraction of saline water and desalination for
124 domestic use, and recharge of the aquifer using treated wastewater. The results showed that
125 this method is effective to controlling SWI and it results in lower salinity and cost.

126
127 The Nile Delta Aquifer (NDA) in Egypt is subject to severe SWI from the Mediterranean Sea.
128 The NDA is one of the largest aquifers in the world (Sherif et al., 1999). Water resources in
129 NDA are threatened by the increasing population, over-pumping, subsidence, coastal erosion,

130 contamination and salinization. These have led to high risks on water resources in the Nile
131 Delta region and management of water resources has become crucial. Various studies on the
132 NDA have indicated that SWI has been extended to 100 km inland. This aquifer is covered by
133 the Delta and has a huge volume of fresh water. However, the quality of the abstracted water
134 is adversely affected by SWI, upconing and contamination. A number of studies have been
135 directed to identify SWI in the NDA using different numerical models (e.g., Farid (1980 and
136 1985), Amer and Farid (1981), Sherif et al. (1990), Darwish (1994), Amer and Sherif (1999),
137 Sherif and Singh (1997) and Saifelnaser and Sherif (2012)). Abd-Elaty et al., (2014) studied
138 the effect of sea level rise and climatic change on SWI in the NDA using SEAWAT. The study
139 indicated that SLR, reduction in surface water recharge and increasing abstraction rate have
140 strong impact on groundwater level but a combination of these scenarios would have larger
141 impact on the aquifer.

142

143 The current research presents a numerical study to control SWI in the NDA using different
144 well systems to those studied before. Different scenarios are considered in this study
145 including, decreasing pumping rates, recharging the aquifer using treated wastewater,
146 abstraction of brackish water a and combination of these three methods.

147

148 **2. Numerical model**

149

150 In this study, the SEAWAT-V4 code was used to simulate SWI and assess different scenarios
151 for controlling SWI in the Nile delta aquifer. SEAWAT simulates variable-density, transient
152 groundwater flow using MODFLOW and MT3DMS programs. The latest version
153 (SEAWAT-V4) was designed to calculate fluid density, viscosity and heat (Langevin et al.,
154 2007).

155

156 **3. Study area**

157

158 The NDA is one of the largest aquifers in the world (Sherif et al, 2012). This study focuses on
159 the Middle part of the Nile Delta aquifer (MND) which is considered the highest part affected
160 by SWI from the Mediterranean Sea (Figure 1a). The abstraction rate from this area is very
161 high which has resulted in more SWI in the middle part. The MND is located between
162 Latitudes $30^{\circ} 20'$ and $31^{\circ} 50'N$, and longitudes $30^{\circ} 10'$ and $31^{\circ} 35'E$. It is bounded by the
163 Mediterranean Sea in the North, River Nile in the South and the East and the two branches of
164 Nile in the West. It has an area of 9,000 km². The stratigraphy of the study area is divided to
165 Pliocene sediments, an upper fluvial sequence (late Pliocene), and Quaternary (Pleistocene
166 and Holocene) sediments lying over the Pliocene sediments. The sediments constitute variable
167 proportions of sand, clay and gravel with lateral variation and variable thicknesses (Sallouma,
168 1983). The thickness of the Quaternary aquifer varies from 200m in the South to 1000m in the
169 North (RIGW, 1992). The main aquifer in the study area is the Quaternary aquifer. The
170 aquifer is composed of loose sand with gravel related to the old litho-facies (Hefny (1980),
171 Sallouma (1983) and Serag El Din (1989)) as shown Figure 1b (Sakr, 2005).

172

173 **3.1 Aquifer geometry**

174 The SEAWAT code is used to simulate heads and solute transport in the aquifer. The domain
175 (9,000 km²) is divided to 157 rows and 141 columns with square cells of dimensions 1 km²
176 (Figure 2a). The depth was divided into 11 layers with varying depth from 200 m in Cairo to
177 1000 m at the sea. The top layer represents clay cap with depth varying from 20m in the South
178 to 50m in the North. The other layers representing the Quaternary aquifer are divided into
179 equal thickness. The aquifer is semi-confined due to the top clay layer. Two vertical sections
180 are taken in X and Y directions as shown in Figures 2b and 2c.

181

182 **3.2 Boundary conditions**

183 The flow and concentration boundary conditions were allocated to the model with zero
184 constant hydraulic head along the shoreline in the North and a specified head of 16.96 m in
185 the South. The East and West boundaries were assigned according to the water levels in the
186 Damietta and Rosita branches of the River Nile. Burrules Lake was assigned as a water body
187 with zero specified head. A concentration of 40000 mg/l was applied along the coastal line,
188 and 35000 mg/l in Burrules Lake which receives water from some drainages in the Nile Delta.
189 The initial groundwater concentration was set to zero. The initial values of hydraulic
190 parameters of the aquifer, obtained from previous studies, were used as input to the model.

191

192 **3.3 Hydraulic parameters**

193 The hydraulic parameters of the study area for each layer, including specific storage (S_s),
194 specific yield (S_y), hydraulic conductivity (K) and porosity were taken from previous studies
195 (e.g., RIGW, 1980). The conductivity in the vertical and horizontal directions (K_{cv} and K_{ch})
196 for the top layer was taken 2.50 and 25-55 mm/day respectively (RIGW, 1980). The average
197 horizontal hydraulic conductivity (K_h) of the Quaternary aquifer sediments was taken 75
198 m/day (RIGW/IWACO, 1999). The specific storage and the storativity (S) were 0.20 and
199 2.5×10^{-3} respectively (RIGW 1980). The longitudinal and lateral dispersivities were taken as
200 100m and 10m respectively, and the diffusion coefficient (D^*) as 10^{-4} m²/day (Sherif et al.
201 (1988)). The aquifer recharge ranges from 0.25 to 0.80 mm/day (RIGW, 1980).

202

203 Figure 3 presents the annual abstraction rates for the Nile Delta aquifer from 1990 to 2016.
204 The abstraction reached 1.60 and 2.60 BCM in 1980 and 1991 (RIGW, 1992) and 3.02 and
205 4.90 BCM in 1999 and 2003 (Morsy, 2008). Also, the annual abstraction reached 3.50 BCM
206 in 2003 (RIGW, 2003) and 7.00 BCM in 2016 (Molle et al., 2016). Mabrouk (2013) indicated

207 that the abstraction rate of the aquifer increases by 0.10 BCM per year while the rate
208 increased to 0.20 BCM per year from 2003 to 2010. The total number of pumping wells in the
209 study area (MDA) is 946 (Figure 4). The total pumping from the MND aquifer was taken as
210 2.22×10^6 m³/day (810×10^6 m³/year) (Morsy, 2008).

211

212 **4. Model Calibration**

213

214 The model was calibrated using field data collected by RIGW (2008). The solute transport
215 model was calibrated in two steps. The first step was done for the hydraulic head using
216 constant density option (visual MODFLOW) through several trials by changing the hydraulic
217 parameters including hydraulic conductivity which ranges from 0.05 to 0.20 m/day for the
218 clay cap and from 25 to 150 m/day for the Quaternary aquifer. The recharge rate ranged from
219 36 to 400 mm/year. Readings from 22 observation wells in the study area were used for
220 calibration. The field data on hydraulic heads were collected by RIGW in 2008. From the
221 model results the mean and absolute residual mean reached -0.07 and 0.144 m respectively
222 while the root mean square was 0.153 m and normalized RMS was 1.341 %. The calibration
223 target of the model was 10% of the difference between maximum and minimum groundwater
224 heads which is about 1.40 m as shown in Figure 5.

225

226 The second stage was carried out using the variable density option (SEAWAT-V4) for
227 concentration parameters of total dissolved solid (TDS) based on the data collected from field
228 survey for groundwater quality monitoring wells and the measured salinity values based on
229 RIGW (2008). The initial time was applied as 0.001 day using time step of 200 days while the
230 model reached the steady state condition after 1825000 days. As the head convergence
231 criterion, the tolerance of 1×10^{-7} m was assigned using the Preconditioned Conjugate
232 Gradient (PCG) solver, while the the Generalized Conjugate Gradient Solver (GCG) was used

233 for the transport equation with a Courant number of 0.75 to limit the time step size with
234 maximum step size of 365 days using the simulation period of 1825000 days.

235

236 Figure 6a shows the distribution of TDS in the MND with average depth ranging from 675 m
237 in the North to 150 m in the South. The Middle Nile Delta aquifer is subject to severe SWI
238 from the Mediterranean Sea (100 km inland intrusion) as shown in Figure 6b. Form the
239 current model results (base case), the 17,500 mg/l equi-concentration line reached to 82.75
240 km from shoreline and transition zone was 38 km as shown in Figure 6c. These results are
241 consistent with field data of SWI in the Nile delta aquifer presented by Shrief et al (2012) and
242 Nofal et al (2015). Moreover the salt mass balance of the aquifer that reached 4.4572×10^{13} kg.

243

244 **5. Results and discussion**

245

246 Different scenarios were studied to investigate SWI in the MND aquifer and to control SWI as
247 will be discussed in the following sections.

248

249 **5.1 Investigation of SWI in the MND**

250 The SEAWAT code was applied to investigate the intrusion of saline water in the MND
251 aquifer. Different scenarios and their impacts on SWI in the MND are studied including; 0.5
252 m SLR, 50% increase in the abstraction rate in pumping wells (946 wells) and combination of
253 these cases.

254

255 **Scenario 1: Sea level rise**

256 In this scenario the effect of sea level rise is studied, where 0.5 m rise in the sea level is
257 considered. The results show that SLR of 0.5 m would increase saltwater intrusion as shown

258 in Figure 7b and the intrusion of the equi-concentration line 17500 mg/l would reach to 84.25
259 km from the shoreline.

260

261 **Scenario 2: Increasing pumping rates**

262 Population growth leads to increase the abstraction from groundwater aquifers. In this
263 scenario, the model was used to assess the impact of increasing pumping rates by 50%. Figure
264 7b shows that increasing pumping rate by 50% would lead to increase in SWI with the equi-
265 concentration line 17500mg/l reaching to 88 km from the shoreline.

266

267 **Scenario 3: Combination of SLR and increasing abstraction rates**

268 In this scenario, the combination of scenarios 1 and 2 (0.5 m SLR and 50% increase in
269 pumping rates) is simulated. The results show that in this case saltwater intrusion increases as
270 shown in Figure 7b. The intrusion of the equi-concentration line 17500 mg/l reached 89.50
271 km from the shoreline. Figure 7a shows that this scenario could result in higher SWI in the
272 MND aquifer than the previous two scenarios.

273

274 **5.2 Control of saltwater intrusion**

275 The numerical model was used to study the control of SWI in the MND using different
276 scenarios including; decreasing pumping rate from the aquifer, increasing recharge to the
277 aquifer, increasing brackish water abstraction from the aquifer and combination of these
278 scenarios. A number of wells were assigned to the model to represent recharge and abstraction
279 from the aquifer. The distribution of 111 recharge wells, 159 abstraction wells and
280 combination are shown in Figure 8a, b and c. The distribution of 946 pumping wells is shown
281 in Figure 4. The results of the four scenarios are presented below and compared with the new
282 position of SWI in the aquifer presented in scenario 3.

283

284 **Scenario 4: Decreasing pumping**

285 It was shown that increasing the pumping rates would increase the intrusion length (scenario
286 2). To prevent such increase in SWI, the pumping rate from 946 wells in the MND is reduced
287 by 25% and 50% of the total pumping rate (2.22×10^6 m³/day). The results show that the equi-
288 concentration line 17500mg/l would reach to 86.50 and 84 km from the shoreline due to
289 reduction of 25% and 50% respectively (Figures 9a and 9b). The relation between intrusion
290 length (X_T) and reduction of pumping rates from wells can be seen in Figure 9c. The results
291 reveal that decreasing pumping rate would lead to decrease in seawater intrusion. The relation
292 can be used to determine the intrusion length (X_T) knowing the pumping rate.

293

294 **Scenario 5: Increasing recharge**

295 Another scenario is considered to control SWI by recharging treated wastewater to the aquifer
296 using wells. 111 wells (150 m deep) in three rows are considered with spacing of 2 km and 5
297 km distance between rows. This water has TDS of 200 mg/l. It is assumed to be taken from
298 treated wastewater treatment. The recharge rate is taken as 125% and 150% of the pumping
299 rate. The results show that intrusion would reach 79 and 72 km from shoreline for the equi-
300 concentration 17500mg/l for 125% and 150% increase in recharge respectively (Figures 10a
301 and 10b). Figure 10c shows the relation between (X_T) and increasing recharge rate. The results
302 indicate that increasing recharge rate leads to decrease in seawater intrusion. The relation can
303 be used to calculate intrusion length (X_T) knowing the recharge rate.

304

305 **Scenario 6: Abstraction of brackish water and disposal to the sea**

306 This scenario is used to study controlling SWI using 159 abstraction wells (350 m deep)
307 installed in three rows in the north of the region with 2 km spacing and 10 km distance

308 between the rows. The abstraction rate is taken as 125% and 150% of the pumping rate.
309 Figures 11a and b show that SWI would reach 87 and 84.25 km from the shoreline for equi-
310 concentration line 17500 for 125% and 150% of abstraction rates respectively. A relation
311 between the intrusion length (X_T) and abstraction rate is shown in Figure 11c. The relation can
312 be used to calculate the intrusion length (X_T) as a function of the abstraction rate. The results
313 indicate that increasing abstraction of brackish water leads to decrease in seawater intrusion.

314

315 **Scenario 7: Combination of scenarios 4, 5 and 6**

316 In this scenario, a combination of the above three scenarios including reducing abstraction
317 rate (scenarios 4), increasing recharge using treated wastewater (scenarios 5) and abstracting
318 brackish water (scenarios 6) is applied to control SWI. The results show that the intrusion
319 could reach to 78 and 70.50 km from shoreline for equi-concentration line 17500mg/l for the
320 cases of 125% and 150% change as shown in Figure 12a and 12b. A relation between
321 intrusion length (X_T) and combined scenario is shown in Figure 12c. The results show that
322 decreasing abstraction of freshwater, increasing recharge and increasing the abstraction of
323 brackish water lead to higher retardation of seawater intrusion. The relation can be used to
324 determine the intrusion length (X_T) as a function of the recharge and abstraction rates.

325

326 The intrusion length in the MND aquifer due to different scenarios is summarized in Table 1
327 which shows a comparison between the different scenarios. The intrusion length reached
328 82.75 km under the current conditions (base case) which increased to 89.50 km (scenario 3)
329 when 150% increase in pumping rate and sea level rise of 50 cm were considered. Different
330 scenarios were considered to manage saltwater intrusion in scenario 3. The first three single
331 well systems used in scenarios 4, 5 and 6 to control SWI are capable of reducing the intrusion
332 and moving the transition zone towards the sea. Decreasing pumping rate by 50% (scenario 4)

333 could move the equi-concentration line 175000mg/l back 5.50 km (6.20%) towards the sea.
334 Increasing recharge by 150% (scenario 5) could move the equi-concentration line 175000mg/l
335 back 17.75 km (19.5%) towards the sea. Increasing abstraction of brackish water from the
336 aquifer by 150% (scenario 6) would move the equi-concentration line 175000mg/l back 5.25
337 km (5.90%) toward the sea. However, combination of these three systems (decreasing
338 pumping, increasing recharge and increasing abstraction of brackish water by 150%) which
339 was presented in scenario 7, could move the equi-concentration line 175000mg/l back 19 km
340 (21.30%) towards. Scenario 7 is considered the best scenario as it gives the highest retardation
341 to SWI. In addition to providing higher retardation for SWI, scenario 7 also has some other
342 advantages including: (i) recharging the aquifer using treated wastewater can help in
343 sustainable management of water resources; and (ii) abstracting brackish water that can be
344 desalinated using e.g., solar-powered desalination plants, can provide additional source of
345 freshwater to meet part of water demand that in turn can help to reduce abstraction of
346 groundwater. This method can control SWI and protect the MND aquifer from deterioration.

347

348 **6. Conclusion**

349

350 Saltwater intrusion (SWI) threatens many coastal aquifers in the world. Increasing rates of
351 abstraction of groundwater, climate change and sea level rise have increased SWI. Saltwater
352 intrusion control is crucial for protection of groundwater storage in coastal areas. The Nile
353 Delta aquifer contains a large quantity of freshwater (400BCM) but it is subject to severe SWI
354 from the Mediterranean Sea. This paper presented a numerical study to simulate SWI in the
355 middle Nile Delta (MND) aquifer. The impact of increasing pumping from the aquifer, sea
356 level rise and combination of the two on SWI was studied, which resulted in increased
357 intrusion by 1.50, 5.25 and 6.75 km respectively. Different scenarios were considered to
358 control saltwater intrusion including; decreasing rate of abstraction from the aquifer,

359 increasing recharge using treated waste water, abstraction of brackish water for desalination
360 and combination of these scenarios. The results indicated that decreasing pumping, increasing
361 recharge and increasing brackish water abstraction (scenarios 4, 5 and 6) can retard saltwater
362 intrusion 6.20%, 19.50% and 5.90% respectively. However, combination of scenarios 4, 5 and
363 6 (scenario 7) can lead to further retardation of saltwater intrusion (21.30%).

364

365 **Data Availability**

366 All data, models, and code generated or used during the study appear in the submitted article

367

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517 Table 1: Intrusion Characteristic in the MND aquifer for different scenarios
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Case	Scenario No.	Scenario description	Intrusion for Iso-concentration line 35		
			Length (Km)	Difference from base case (km)	Percentage (%)
Base case	-	Current situation	82.75	-	-
Different scenarios of SWI investigation	I-1	0.5 Sea level rise	84.25	-1.5	-1.81
	I-2	50% Increasing pumping	88	- 5.25	- 6.34
	I-3	Combination of 1 and 2	89.50	- 6.75	- 8.16
Different scenarios SWI management	M-4	-25% wells pumping rates	86.50	- 3.75	- 4.53
		-50% wells pumping rates	84	- 1.25	- 1.51
	M-5	+25% wells recharge	79	+ 3.75	+ 4.53
		+50% wells recharge	72	+ 10.75	+ 12.99
	M-6	+25% wells abstraction	87	- 4.25	- 5.14
		+50% wells abstraction	84.25	- 1.5	- 1.81
	M-7	-25 and +25% combination of 4, 5 and 6	78	+ 4.75	+ 5.74
-50 and +50% combination of 4, 5 and 6		70.50	+ 12.25	+ 14.80	

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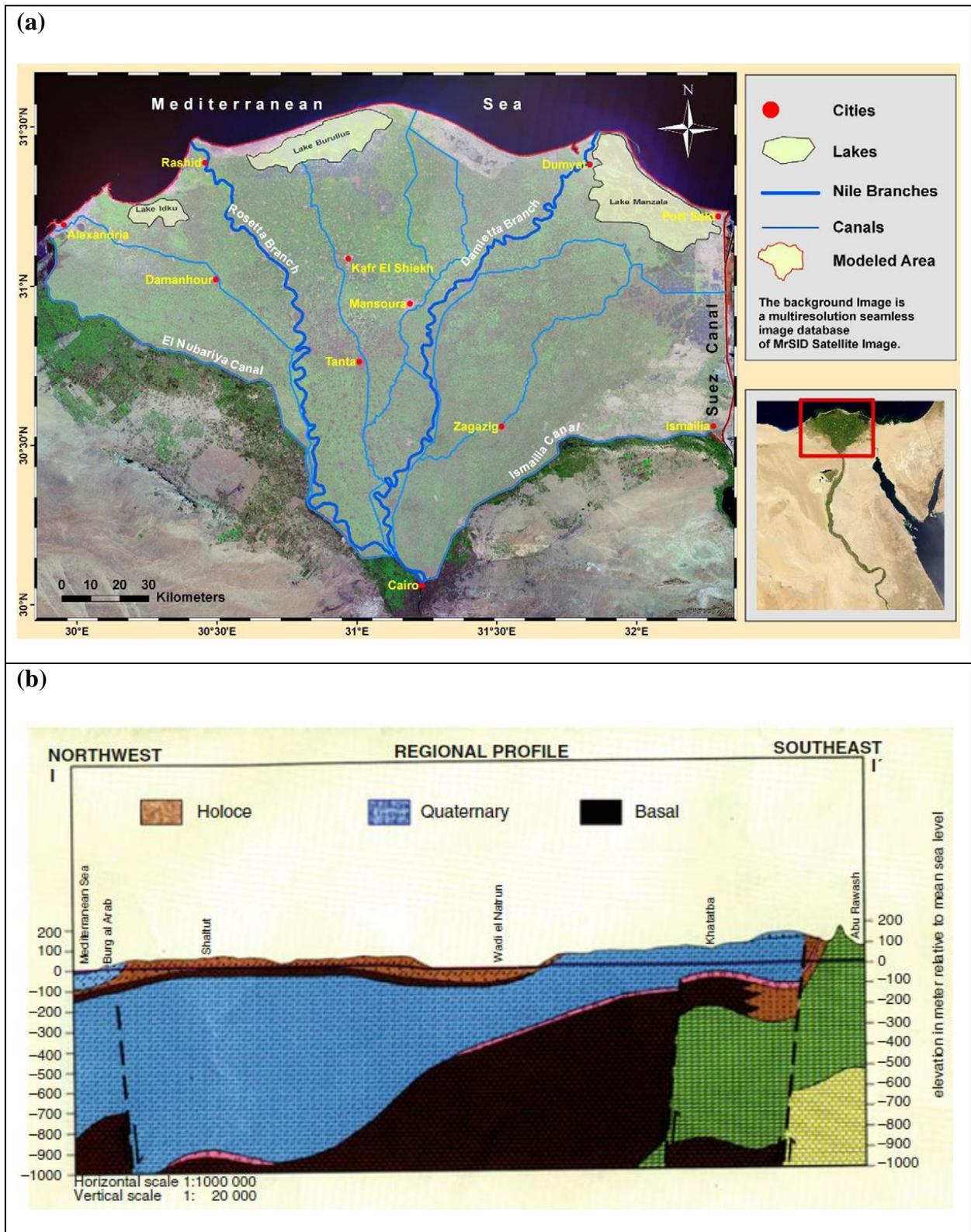
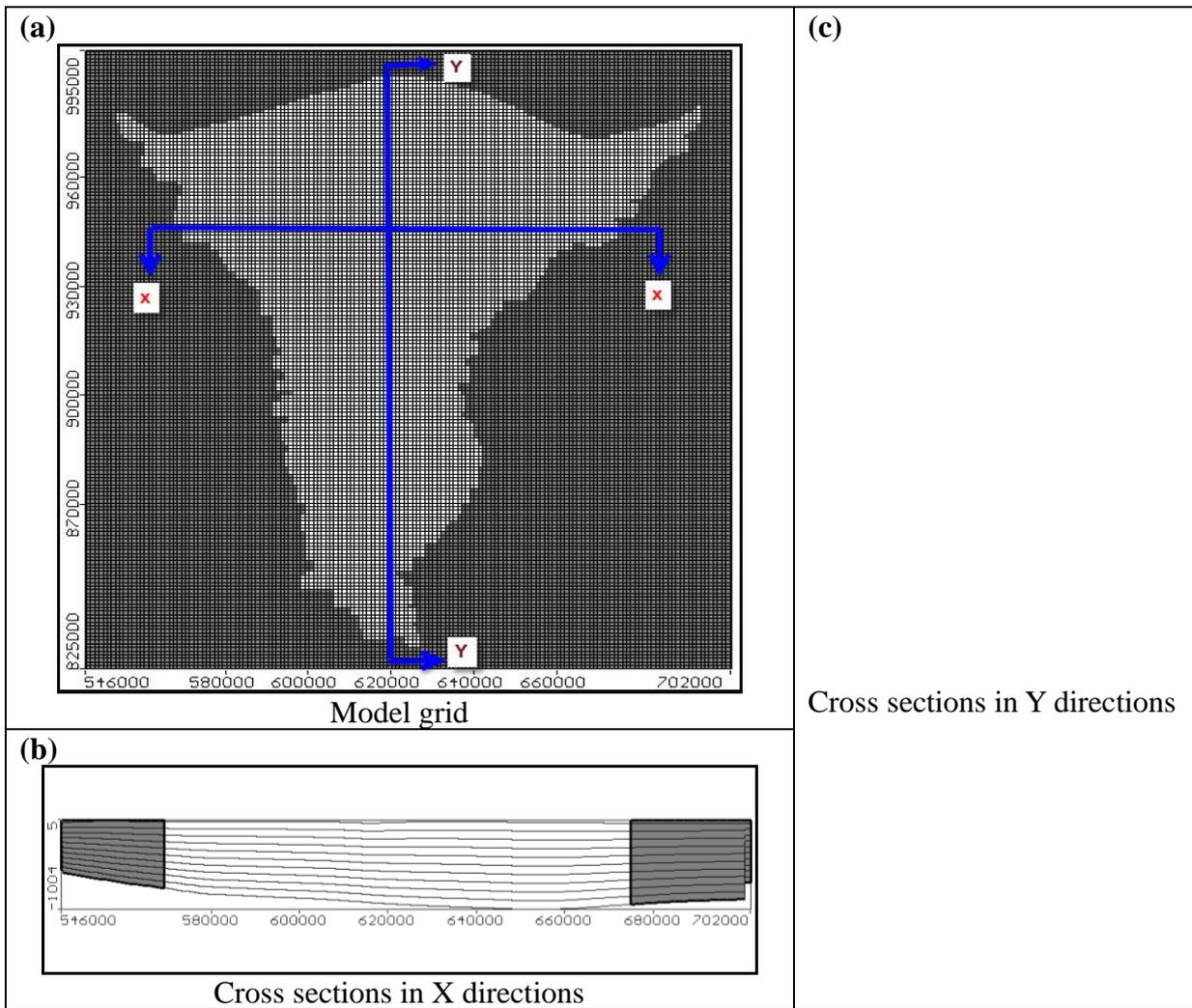


Figure 1. The Nile Delta aquifer (a) Location map (Sherif et al, 2012) and (b)

Hydrogeological cross section (Sakr, 2005)



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527 Figure 2: Model grid and cross sections in X and Y directions in the MND aquifer

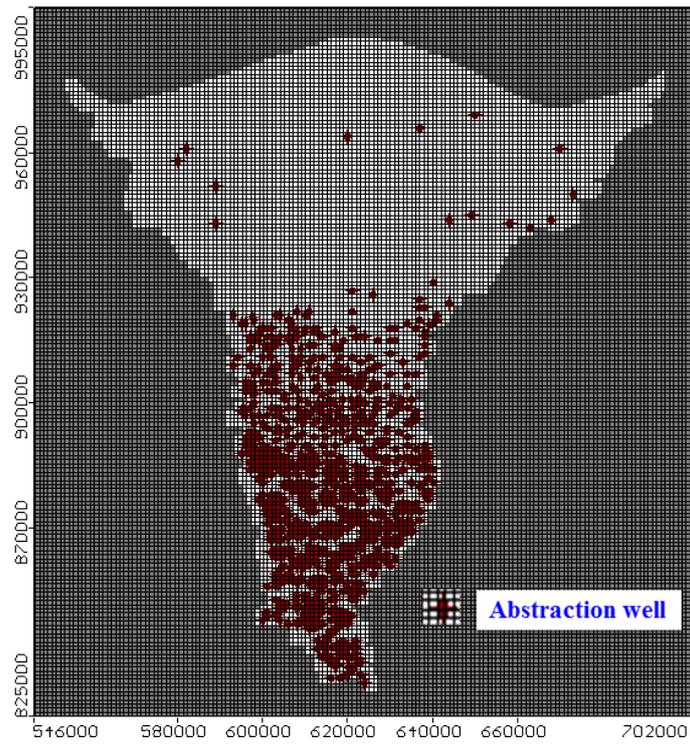
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Figure 3: Annual abstraction rates of Nile delta aquifer

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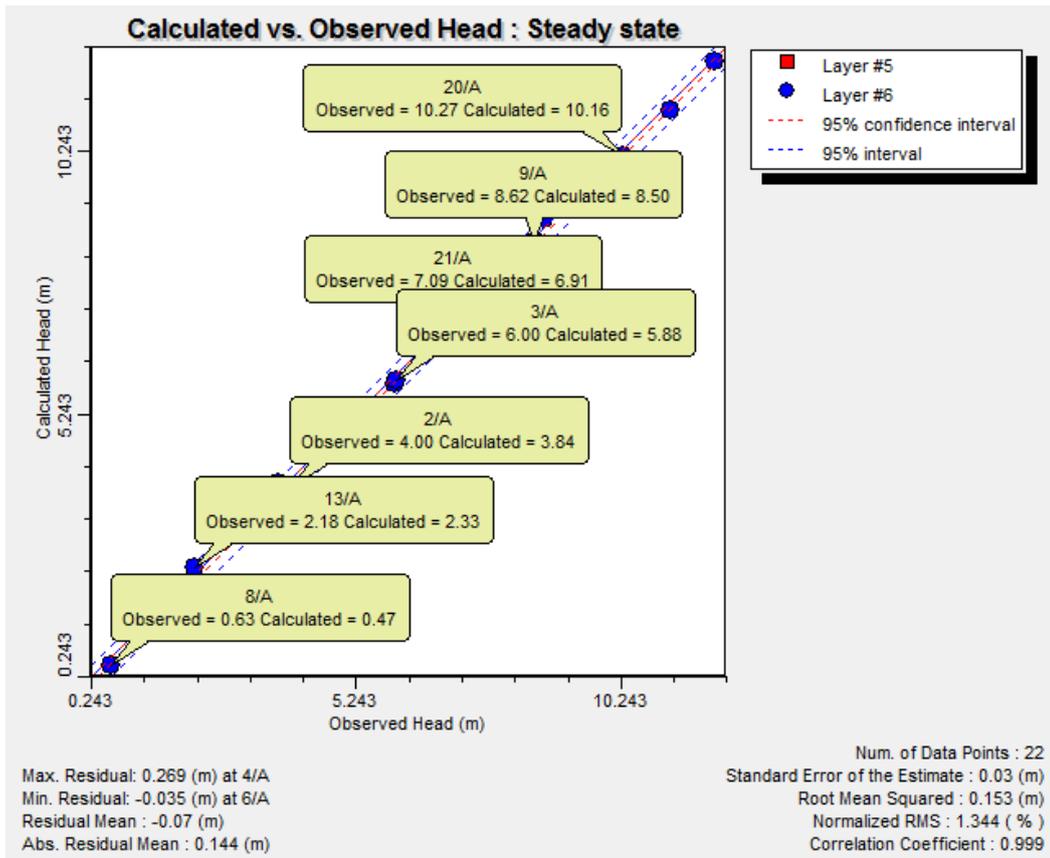
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Figure 4: Distribution of abstraction wells in the MND aquifer in year 2008

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Figure 5: Calculated and observed groundwater heads for the MND aquifer

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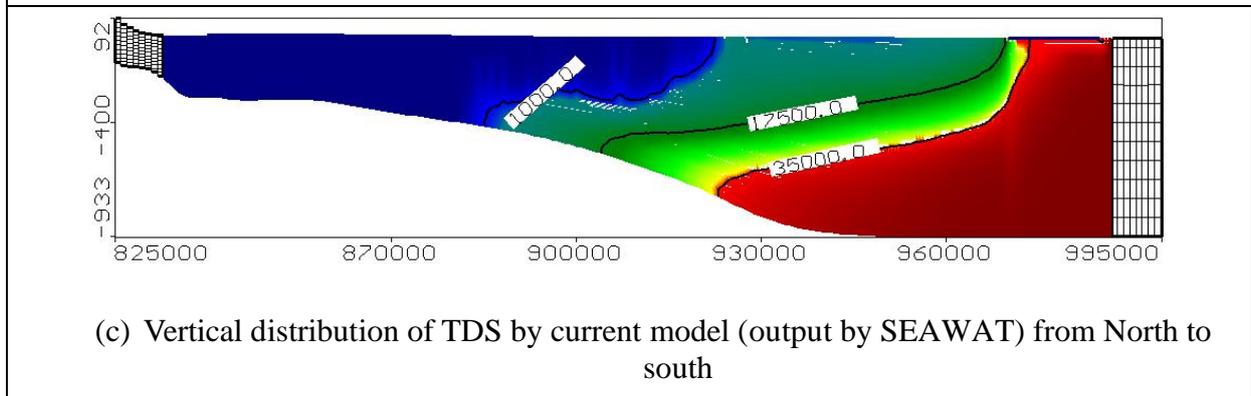
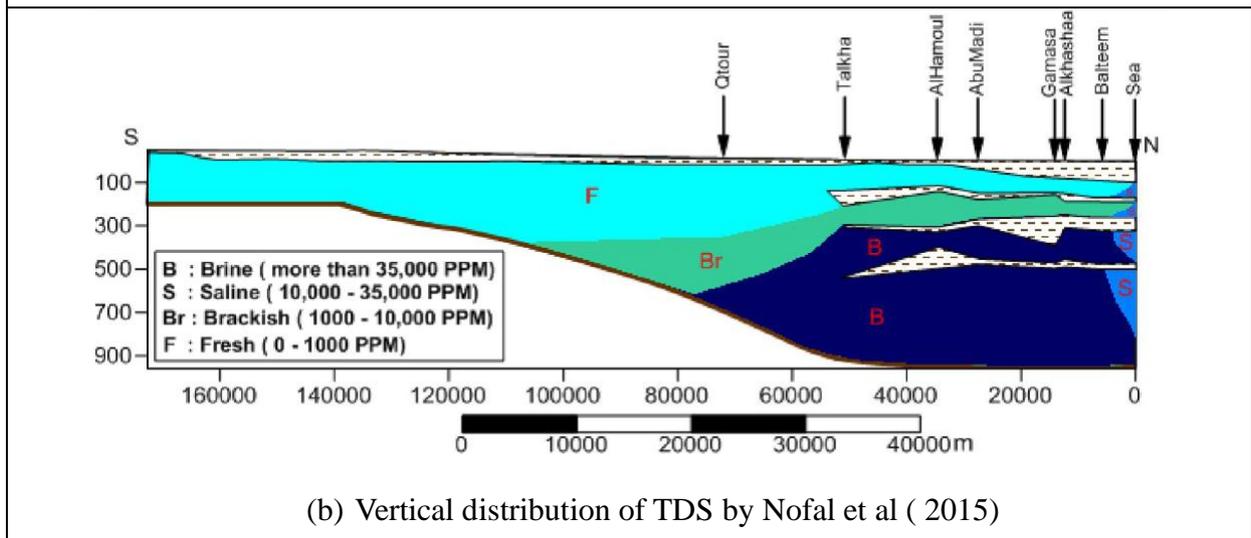
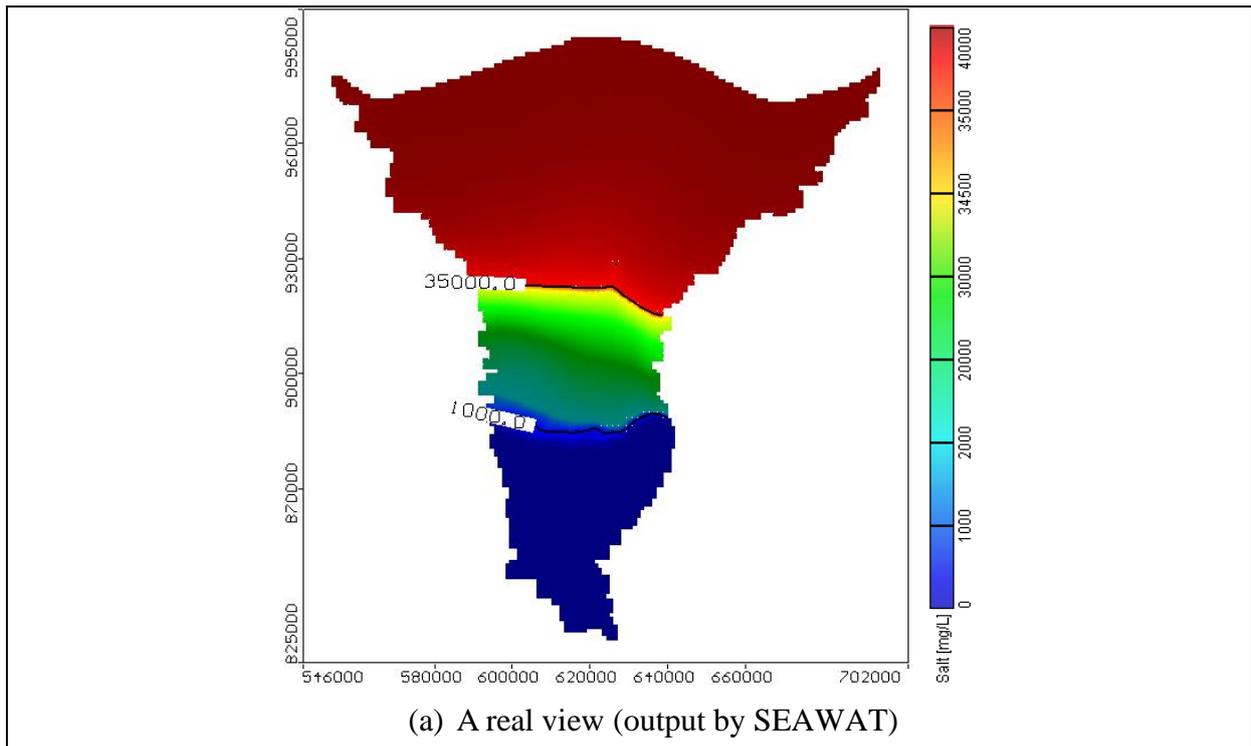
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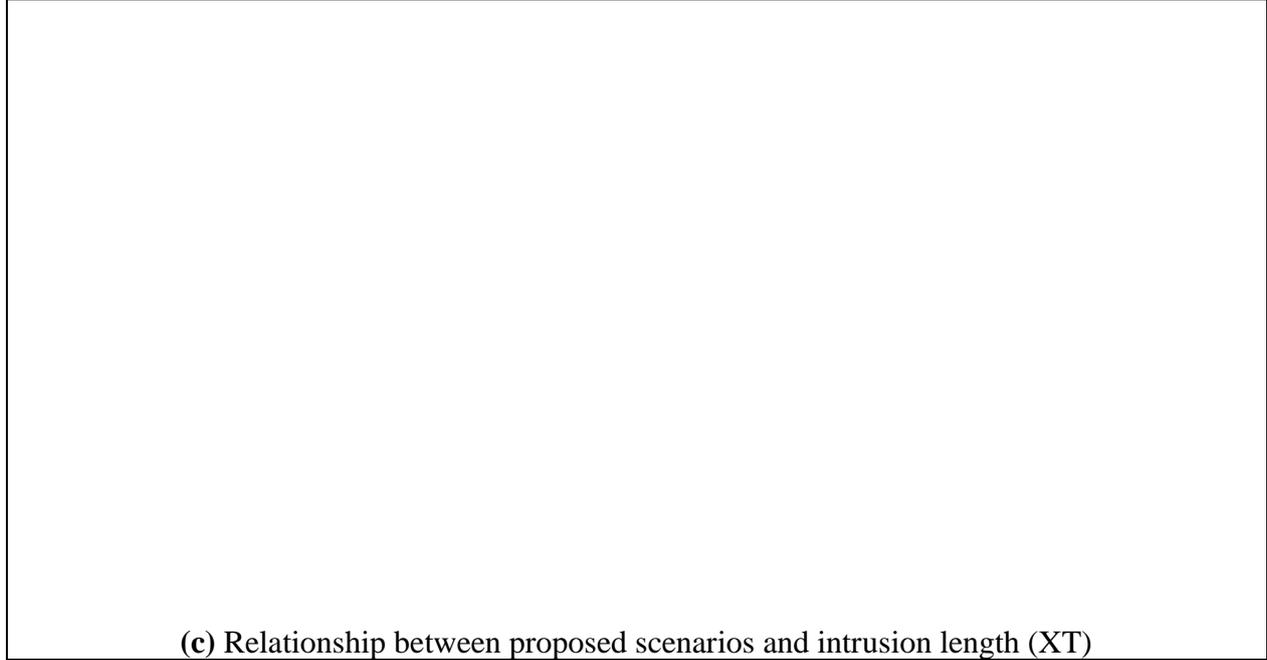
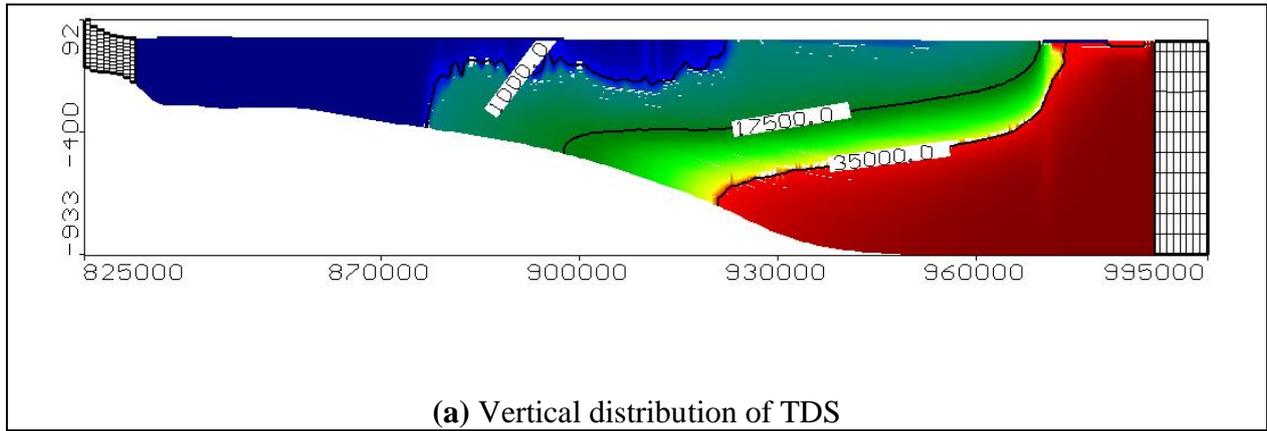
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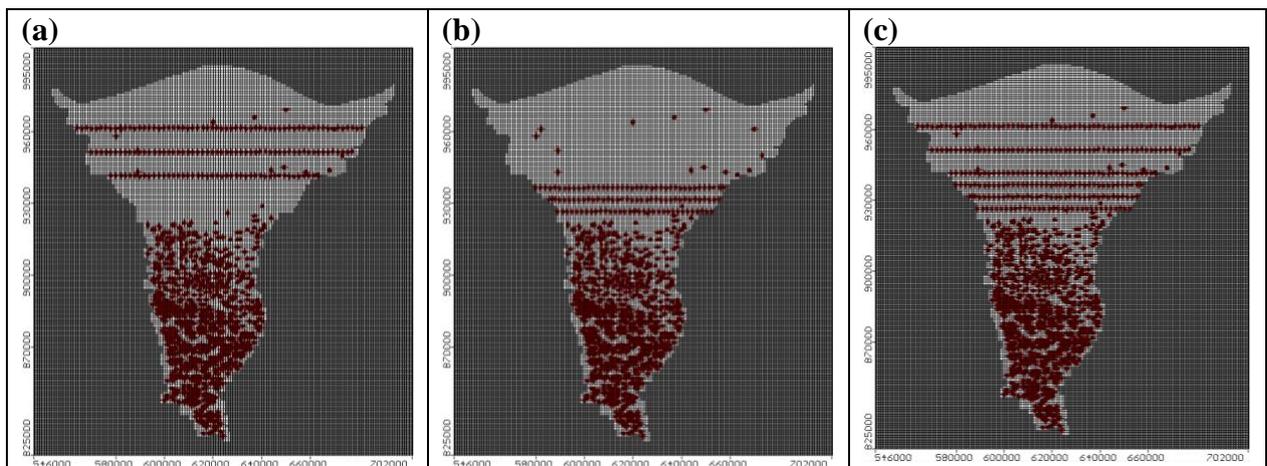
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Figure 6: TDS investigation results in the MNDA for base case in year 2008



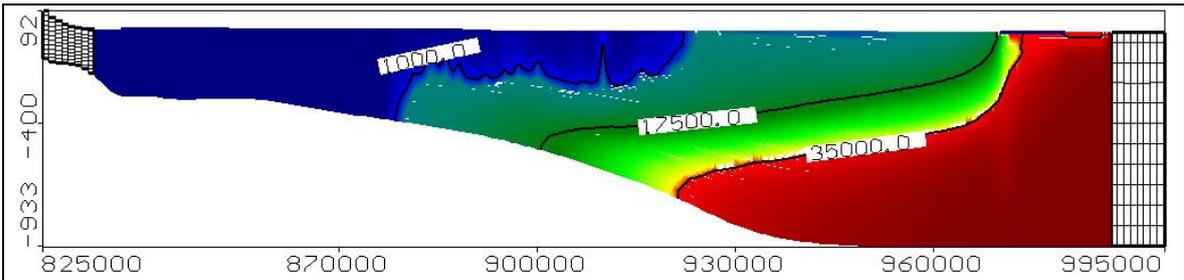
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Figure 7: TDS investigation results for combination of 50 cm SLR and 50% over-pumping in the MND (Scenario.1, 2 and 3)

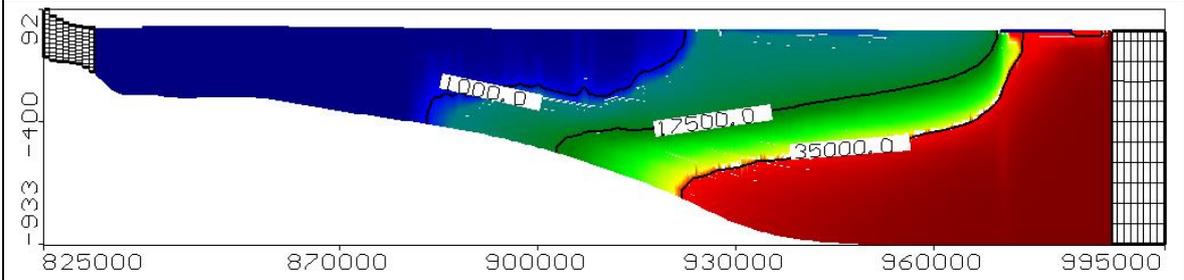


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Figure 8: Areal distribution of (a) abstraction wells, (b) recharge wells, and (c) combination of abstraction and recharge wells in the MND aquifer



(a) Vertical distribution of TDS at -25% in well abstraction

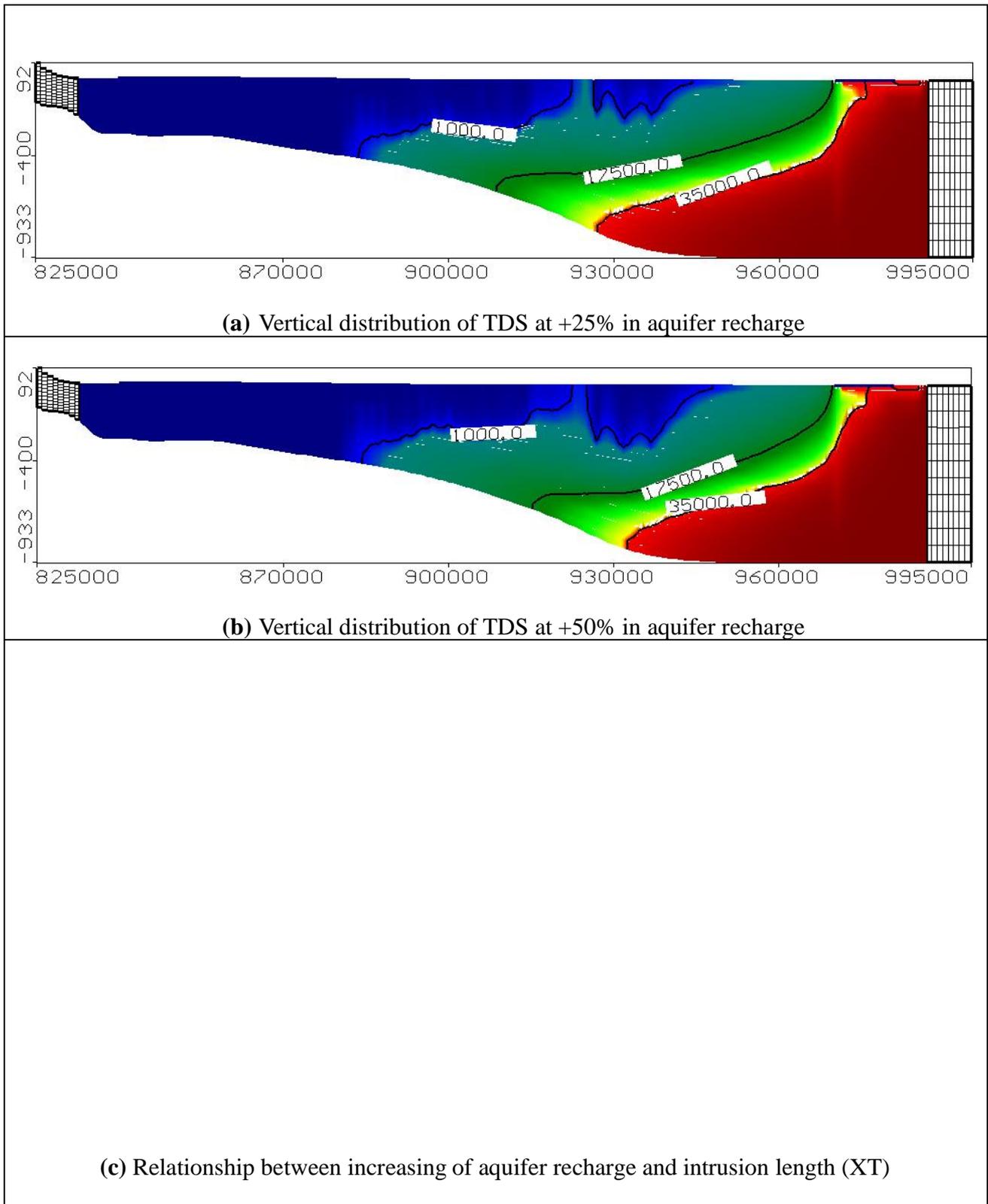


(b) Vertical distribution of TDS at -50% in well abstraction

(c) Relationship between reduction of pumping rates and intrusion length (XT)

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Figure 9: TDS management results for reducing pumping rates in the MND (Scenario.4)

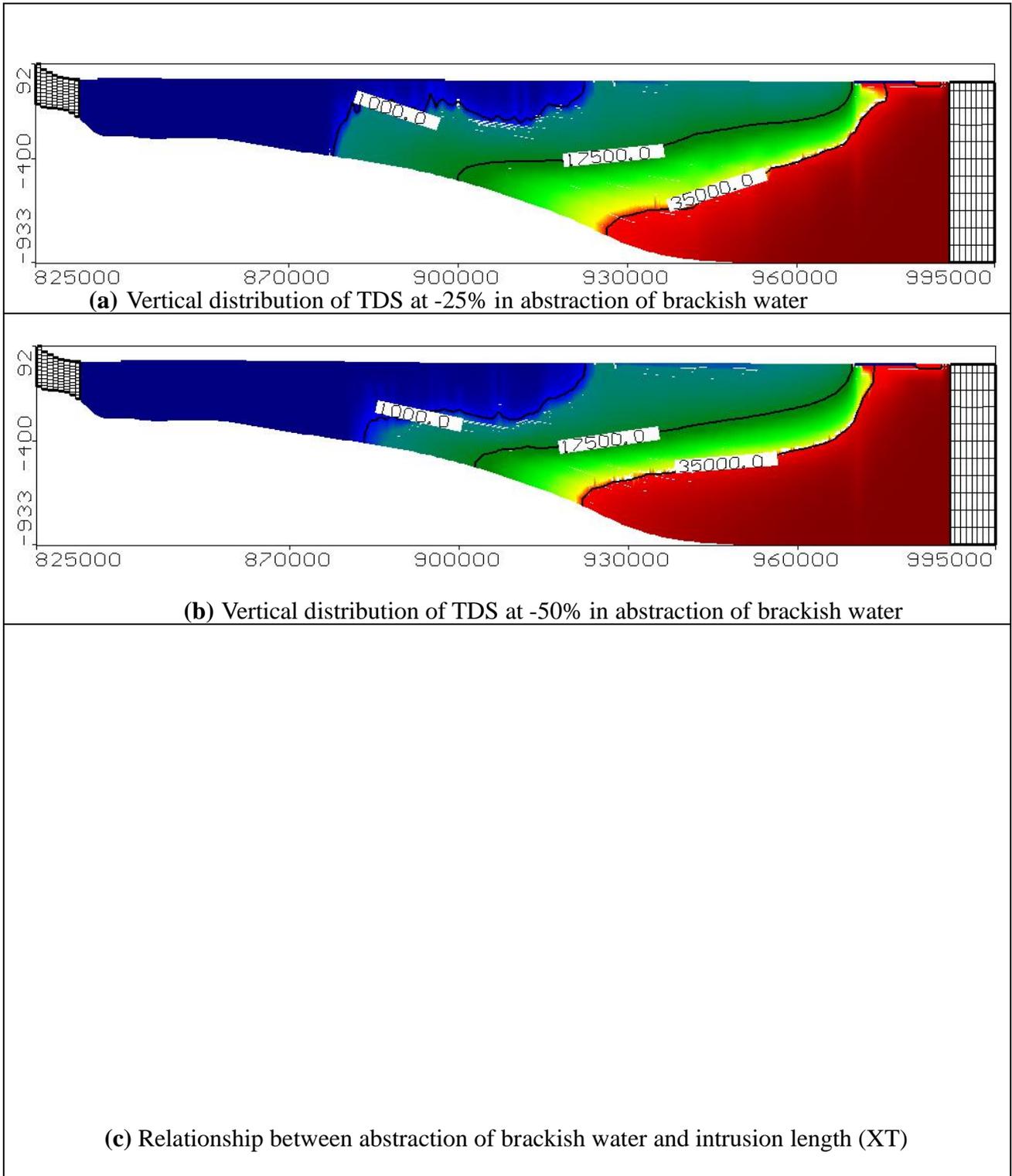


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578 Figure 10: TDS management results for increasing aquifer recharge in the MND (Scenario.5)

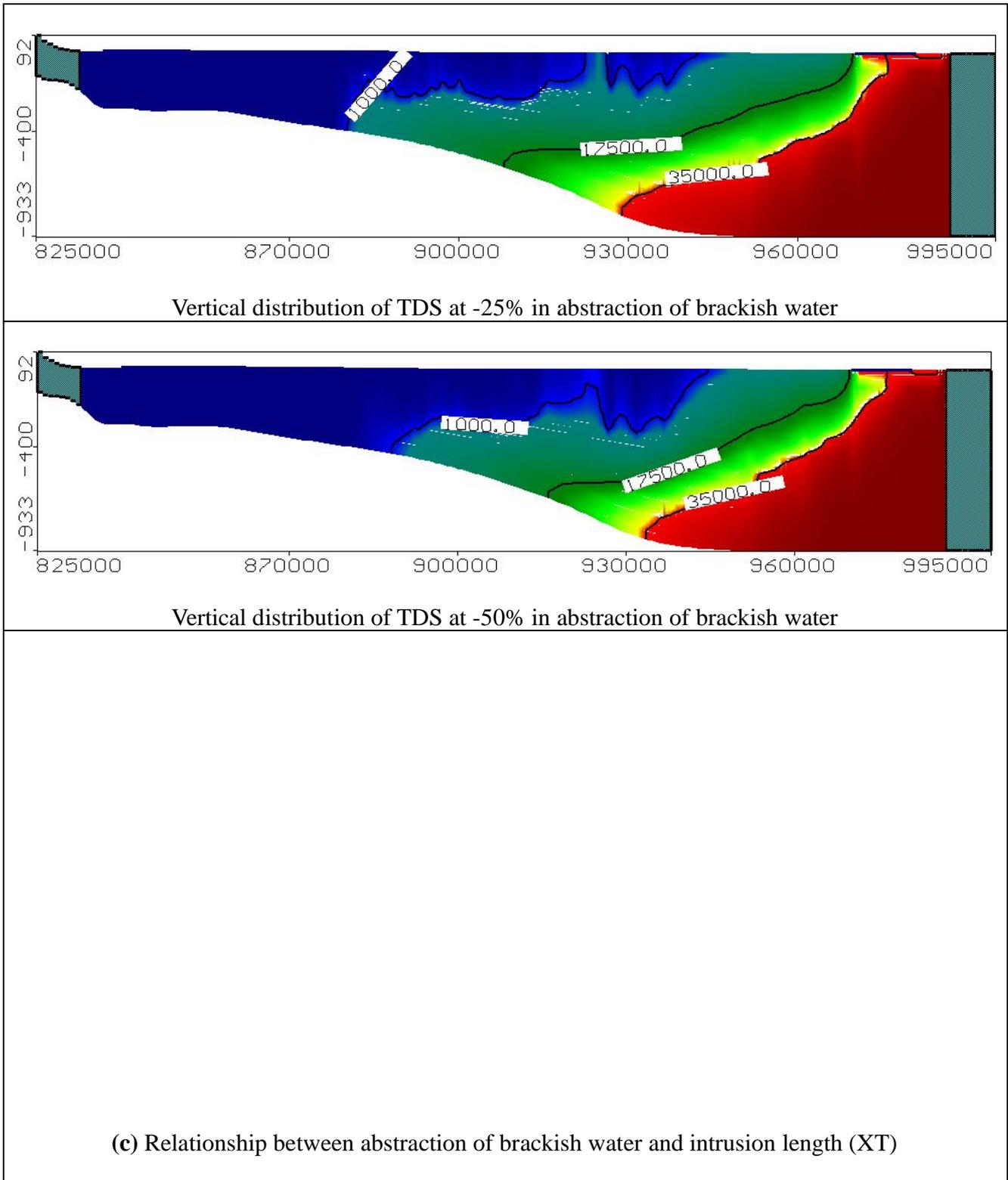
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Figure 11: TDS management results for increasing abstraction of brackish water in the MND (Scenario.6)



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Figure 12: TDS management results for the combination of scenarios 4, 5 and 6 at changes in the MND (Scenario.7)