

1 **Prebiotic Chemistry and Atmospheric Warming of Early Earth By An Active**
2 **Young Sun**

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11 Nitrogen is a critical ingredient of complex biological molecules [1]. Molecular nitrogen,
12 however, which was outgassed into the Earth's early atmosphere [2], is relatively chemically
13 inert and nitrogen fixation into more chemically reactive compounds requires high temperatures.
14 Possible mechanisms of nitrogen fixation include lightning, atmospheric shock heating by
15 meteorites, and solar ultraviolet radiation [3,4]. Here we show that nitrogen fixation in the early
16 terrestrial atmosphere can be explained by frequent and powerful coronal mass ejection events
17 from the young Sun - so-called superflares. Using magnetohydrodynamic (MHD) simulations
18 constrained by Kepler space telescope observations, we find that successive superflare ejections
19 produce shocks that accelerate energetic particles, which would have compressed the early
20 Earth's magnetosphere. The resulting extended polar cap openings provide pathways for
21 energetic particles to penetrate into the atmosphere and, according to our atmospheric chemistry
22 simulations, initiate reactions converting molecular nitrogen, carbon dioxide and methane to the
23 potent greenhouse gas nitrous oxide as well as hydrogen cyanide, an essential compound for life.

24 Furthermore, the destruction of N₂, CO₂ and CH₄ suggests that these greenhouse gases cannot
25 explain the stability of liquid water on the early Earth. Instead, we propose that the efficient
26 formation of nitrous oxide could explain a warm early Earth.

27 **MAIN TEXT**

28 Here we develop a new concept of the rise of prebiotic chemistry on early Earth that suggests
29 abiotic nitrogen fixation mediated by the energy flux from paleo solar eruptive events. The flare
30 statistics of *Kepler* data suggests that the frequency of occurrence of superflares with energies >
31 5×10^{34} erg observed on G-type dwarfs follows a power-law distribution with spectral index
32 between ($\alpha = -2.0$), which is comparable to those observed on dMe stars and the Sun [5,6]. If the
33 occurrence rate of superflares on young solar-like ~ 0.1 events/day [6], then, the frequency of
34 super Carrington-type flare events with $E \sim 10^{33}$ ergs on the early Sun (≤ 0.5 Gyr) is expected to
35 be ~ 250 events per day! Current data suggest that powerful solar flares (over X5 type) are
36 usually associated with fast (≥ 1000 km s⁻¹) wide ($\theta > 100^\circ$) coronal mass ejections (CMEs) and
37 high-fluence solar energetic particle (SEP) events with kinetic energies up to 10^{33} ergs [7-9].
38 Tree ring data have recently provided evidence in favor of past superflares from the Sun [10,11].
39 Their energy is a factor of 2-3 greater than that suggested for the famous Carrington-type CME
40 event [12]. Recent direct measurements of surface longitudinal magnetic fields on young solar-
41 type stars imply that our young Sun had generated at least 10 times greater magnetic flux than
42 that observed in the current Sun [13]. The stronger magnetic flux produces frequent and
43 energetic flares, fast and wide coned CMEs and associated energetic SEP events with energies
44 up to 10^{36} ergs. Our calculations suggest that the probability of CME striking the Earth is about
45 5% [14]. In the “perfect” magnetospheric storm, when the incoming cloud magnetic field, B_z
46 component is sheared with respect to the Earth's magnetic field, the frequency of CME impacts

47 is > 1 event per day! To model a CME event and its effects on the Earth's magnetosphere, we
48 used the Space Weather Modeling Framework (SWMF) available through the Community
49 Coordinated Modeling Center (CCMC) (see Supplementary Material). We assumed a steady
50 state paleo solar wind at 0.7 Gyr with the mass loss rate of $1.7 \times 10^{-12} M_{\text{sun}}/\text{yr}$ and the wind speed
51 of 700 km/s as obtained from the 3D MHD young Sun's wind model [15] and a Carrington-type
52 CME cloud propagating at the radial speed of 1800 km/s with the total energy of 2×10^{33} erg
53 [12]. Figure 1 presents a 2D map of the steady-state plasma density superimposed by magnetic
54 field lines for the magnetospheric configuration in the $Y=0$ plane corresponding to the initial 30
55 minutes of the simulations, when the Earth's magnetosphere was driven only by dynamic
56 pressure from the paleo-solar wind. The left panel of Figure 1 shows the steady state paleo solar
57 wind compresses the Earth's magnetosphere to $\sim 9 R_E$. The right panel of Figure 1 shows the
58 state of the magnetosphere two hours later when the CME cloud hits the Earth's magnetosphere
59 (also see the movies in Supplementary Material). At this time, the solar wind dynamic pressure
60 and the magnetic reconnection between the southward directed CME's cloud magnetic field and
61 northward Earth's dipole field pushing the dayside magnetosphere earthward reducing the
62 magnetopause stand-off distance from 9 to ~ 1.5 Earth's radii. The CME drives large field aligned
63 current distributions and produces significant disturbance of the magnetospheric field shifting the
64 boundary of the open-closed field shifts to 36° in latitude and producing a polar cap opening to
65 70% of the Earth's dipole magnetic field. In the current version, we used the dipole magnetic
66 field of the current Earth, however paleomagnetic studies of the Earth's ancient rocks suggest
67 that the field was weaker [16]. This suggests that the fraction of the open field used in our model
68 represent only a lower bound. Energetic particles accelerated in shocks driven by successive
69 flare/CME events (see for example [17]) can then efficiently penetrate the early terrestrial

70 atmosphere through the expanded polar cap region.

71 We applied the Aeroplanet model [18] to simulate the atmospheric chemistry of the
72 nitrogen-dominated (80% N₂, 20% CO₂ and 0.03% CH₄) primitive Earth's atmosphere [19]. The
73 upper boundary of the atmosphere at 100 km is exposed to the steady state XUV flux with the
74 spectrum reconstructed for the early Sun at 0.7 Gyr [20] and to energetic protons with the energy
75 flux of 5×10^{11} protons/cm²/MeV at 0.1 MeV with the spectral index of the energy spectrum of -
76 2.15 representative of the Jan 20, 2005 SEP event and the energy range within 1 GeV [21]. The
77 model calculates photoabsorption of the EUV-XUV flux from the early Sun (see Figure 2) and
78 particle (electron and proton) fluxes to compute the corresponding fluxes at the atmospheric
79 altitudes between 200 km and the surface. These fluxes are used to calculate the photo and
80 particle impact ionization/dissociation rates of the atmospheric species producing secondary
81 electrons due to ionization processes. Then, using the photon flux and the photoionization-
82 excitation-dissociation cross-sections, the model calculates the production of ionized and excited
83 state species and as a result, photoelectrons. In our steady-state model of the early Earth's
84 atmosphere, energetic precipitating protons from an SEP event impacted the middle and low
85 atmosphere and produce ionizations, dissociations, dissociative ionizations, and excitations of
86 atmospheric species and as a result, secondary electrons. The model includes 117 neutral
87 chemical reactions. The destruction of N₂ into reactive nitrogen, N(²D) and N(⁴S) and the
88 subsequent destruction of CO₂ and CH₄ produces NO, CO, CH and NH in the polar regions of
89 the atmosphere as shown in Figure 3 (see Supplementary Material).

90 Our model predicts the formation of abundant NO and NH molecules and efficient
91 formation of N₂O through $\text{NO} + \text{NH} \rightarrow \text{N}_2\text{O} + \text{H}$ with the major sink through the reaction $\text{N}_2\text{O} +$
92 $\text{H} \rightarrow \text{OH} + \text{N}_2$ (see the pathway diagram in Figure 3). Photolysis of N₂O via the reaction

93 pathway $\text{N}_2\text{O} + h\nu \rightarrow \text{N}_2 + \text{O}(^1\text{D})$ is not an efficient loss channel for N_2O , because of absorption
94 of solar flux shorter than 2300 Å by CH_4 . Atmospheric N_2O steady-state density reaches a
95 concentration with the mixing ratio of 2 and 20 ppbv at 30 km in the 1 PAL (present atmospheric
96 level) atmosphere with 100% (solid line) and 10% (dashed line) of the maximum photochemical
97 destruction rate, as shown in Figure 4a. The derived value at 100% of the photodestruction rate
98 should be considered as a lower bound, because our model does not account for a number of
99 factors including the eddy diffusion and convection effects, the effects of Rayleigh scattering of
100 solar EUV radiation in the atmosphere and formation of hazes that significantly reduces the
101 photo-destruction rate of nitrous oxide, and therefore increases the production of N_2O . Thus, the
102 model with 10% of the maximum photo-destruction rate probably better represents the density
103 profiles when all factors are accounted for. The steady-state density of N_2O reaches 20 to 3000
104 ppbv in the 2 PAL model with 100% (solid line) and 10% (dashed line) of the maximum
105 photochemical destruction rate, as shown in Figure 4b. The choice of 2 PAL in Figure 4b is
106 consistent with compelling evidence that the atmospheric pressure of early Earth was enhanced
107 by a factor of 2-3 [22]. Another factor affecting the equilibrium mixing ratios of Figure 4 is the
108 representative energy of SEP events, which could be greater than that assumed in the model.
109 Laboratory experiments report the production of nitrogen oxides and N_2O when N_2 - CO_2 mixture
110 that simulates the early Earth atmosphere was exposed by lightning and coroneae discharges [23].
111 Enhanced production of nitrous oxide in the lighting experiments are caused by energetic
112 electrons accelerated in the discharge and UV emission. Other evidence for the role of energetic
113 particles in N_2O production comes from direct observations of its enhancement by 3% associated
114 with thunderstorm events [24].

115 The efficient production of N_2O in our model offers a solution to warming the early

116 Earth. The 0.7 Gyr old Sun was 25-30% fainter than the present-day Sun [25], which would be
117 insufficient to support liquid water on the early Earth contrary to geological evidence of that time
118 [26]. Current models of atmospheric warming offer solutions of this problem, commonly known
119 as Faint Young Sun (FYS) paradox due to a large atmospheric concentration of CO₂, H₂O, CH₄
120 or N₂ and H₂ [27]. However, as our model implies, these molecules will be efficiently dissociated
121 due to photo-collisional processes driven by SEPs from the young Sun, which is consistent with
122 the recent mineralogical data [28]. Instead, the production of CH, NH and NO sets stage for the
123 formation of N₂O, HCN and other N-containing species in the lower parts of the atmosphere.
124 HCN concentration reaches up to tens ppmv in the lower atmosphere. The calculated production
125 rate of HCN in the low atmosphere is driven by the following major reactions: NO + CH →
126 HCN + O, CH₂ + N(⁴S) → HCN + H, CH₃ + N(⁴S) → HCN + H + H, CH + CN → HCN + H,
127 N₂O + CH → HCN + NO. Organic molecules may subsequently rain out into surface reservoirs
128 and engage in higher order chemistry producing more complex organics. For example, further
129 HCN polymerization is known to produce various amino acids, the building blocks of proteins
130 [29]. Production of other types of soluble N-containing species (NH₃, HNO, NO) by particles
131 may have provided a massive dose of nitrogen “fertilizer” to early surface biology on terrestrial
132 planets.

133 Thus, our concept implies that early Sun’s activity provided a window of opportunity for
134 prebiotic life on Earth. The proposed model also redefines the conditions of habitability not just
135 in terms of a “liquid water zone”, but as a biogenic zone (BZ), within which the stellar energy
136 fluxes are high enough to ignite reactive chemistry that produces complex molecules crucial for
137 life. As a by-product, this chemistry forms greenhouse gasses that may efficiently keep the
138 atmosphere warm for liquid water to exist. The model predictions can be tested by observing broad

139 and deep molecular absorption lines of N_2O at $4.5 \mu m$ and $7.9 \mu m$ and HCN absorption features
140 at 3 and $14.3 \mu m$ using James Webb Space Telescope NIRSpec and MIRI observations of
141 primitive terrestrial-type atmospheres around active stars.

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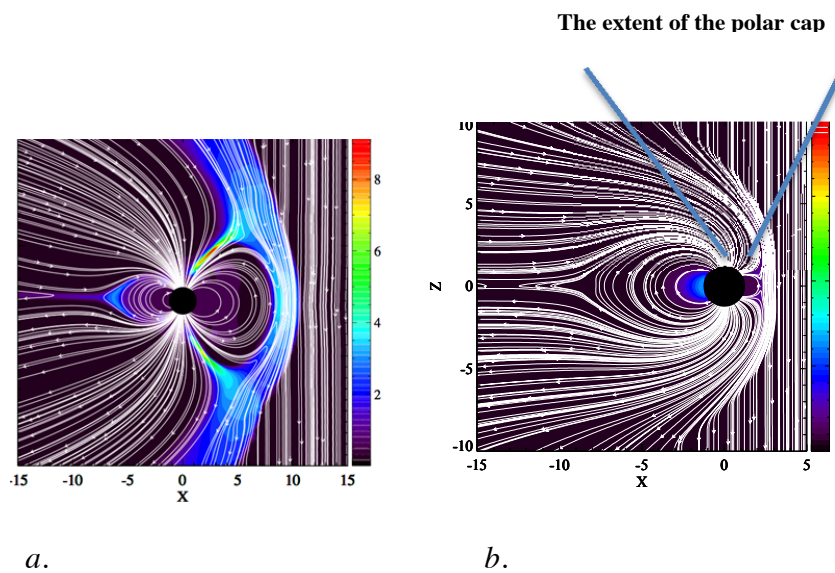
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FIGURES

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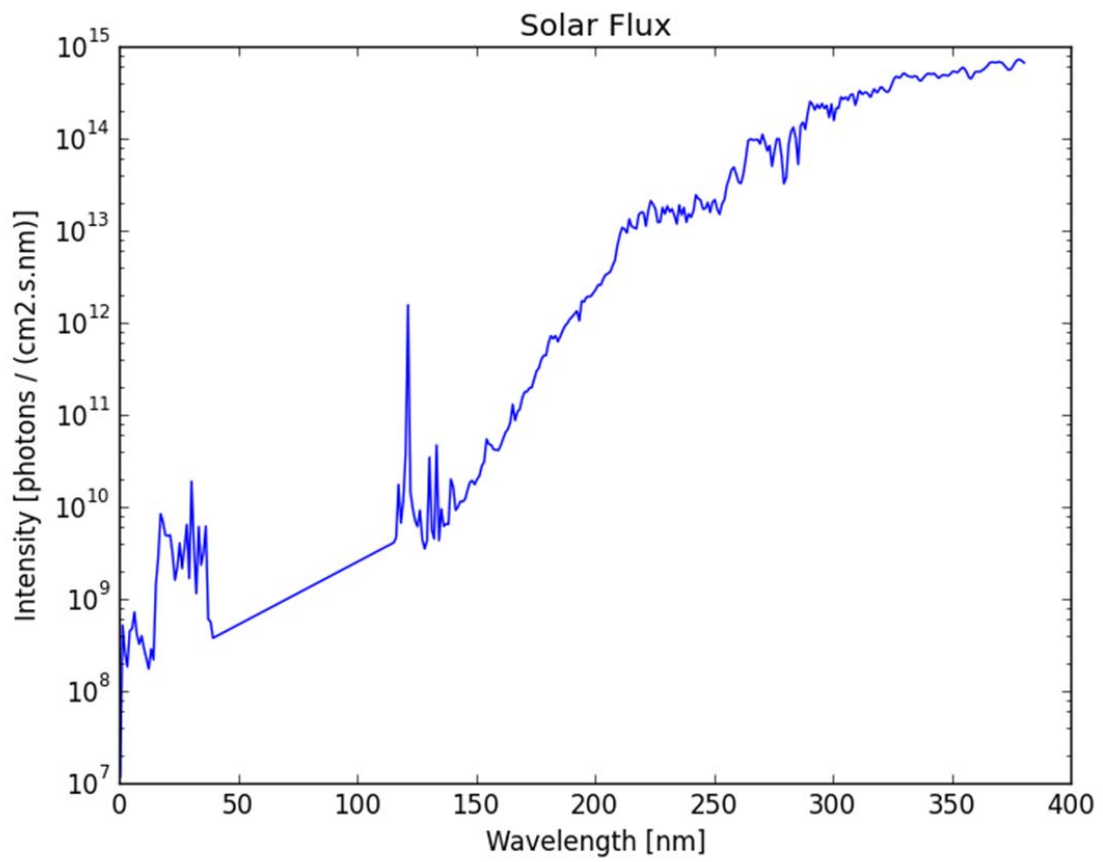
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151 *Figure 1. The initial (panel a) and the final state (panel b) magnetic field lines in white) and the*
152 *plasma pressure (in color) of the Earth's magnetosphere due to the CME event*

153 *(Airapetian et al. 2015)*

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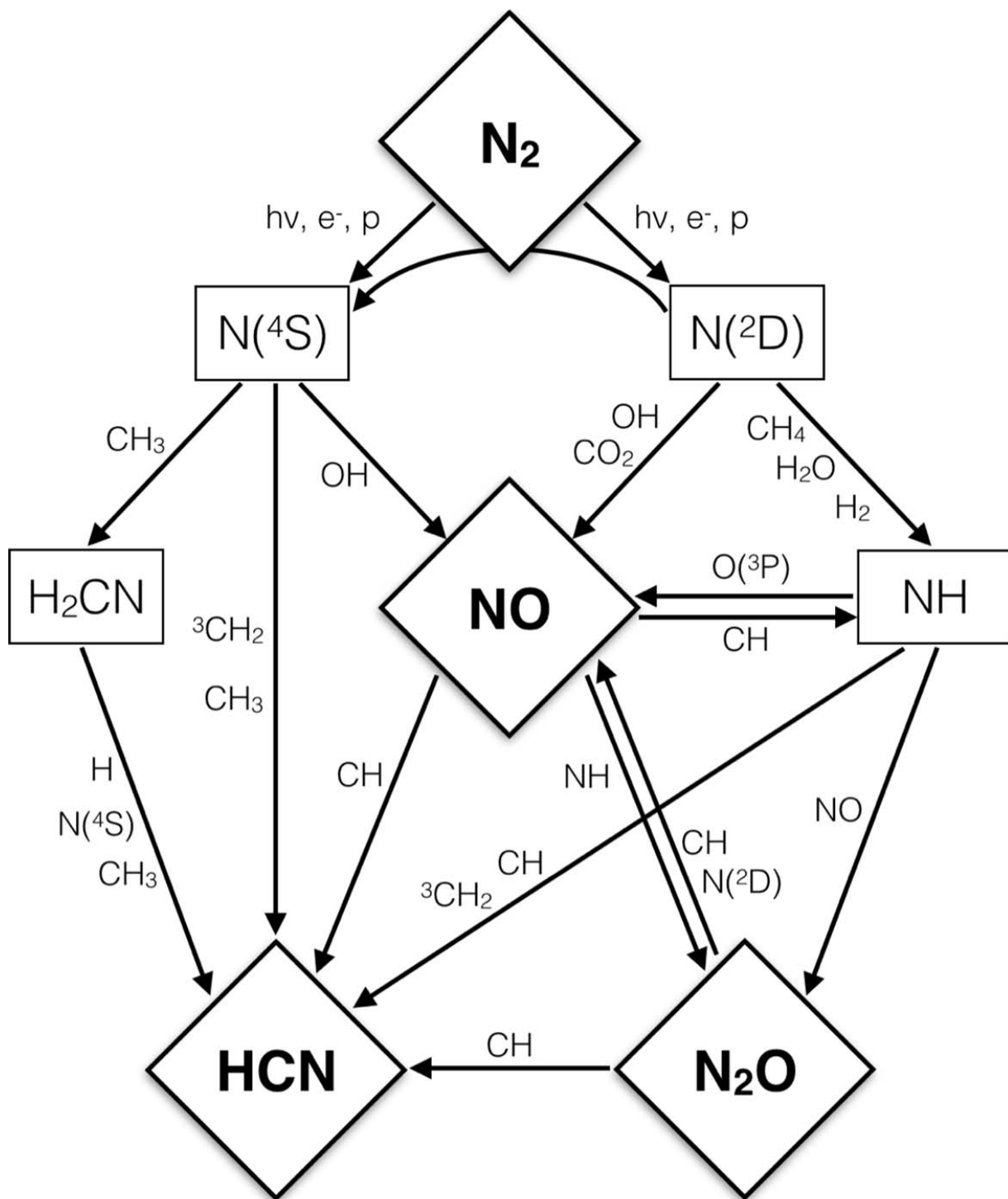
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Figure 2. Spectrum of the young Sun's XUV flux at 0.7 Gyr [20]

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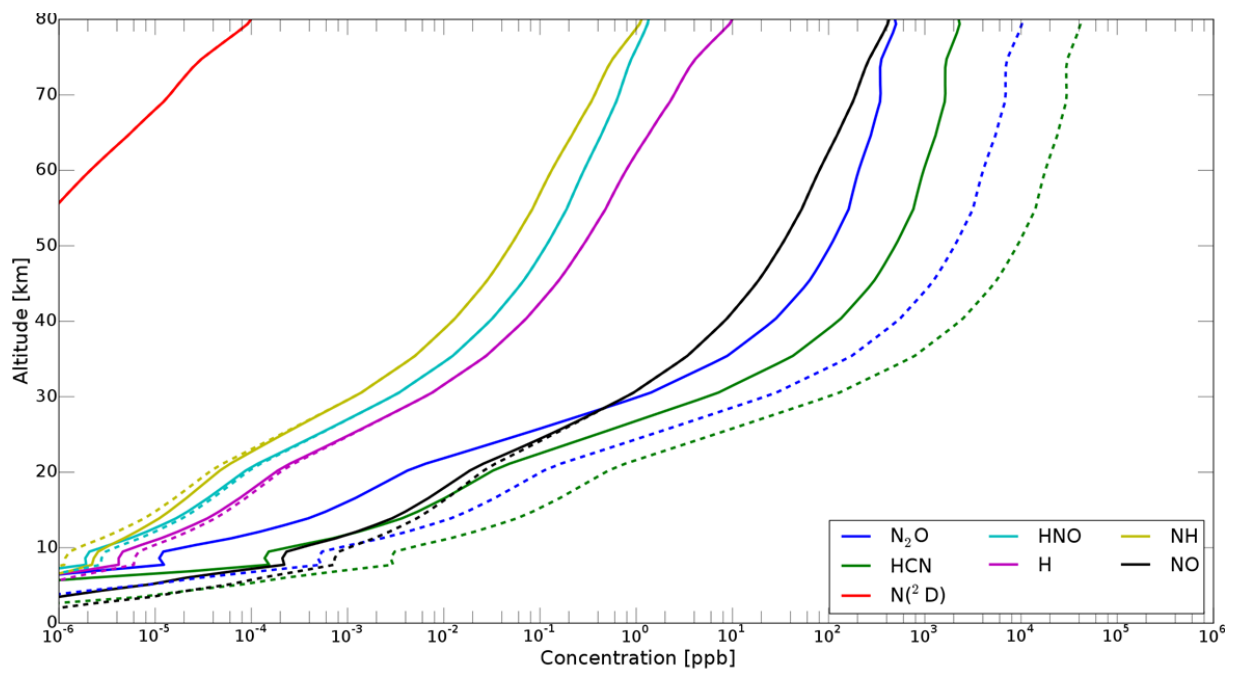
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162 *Figure 3. The pathway diagram of abiotic production of odd nitrogen and nitrogen-bearing*
163 *compounds including nitrous oxide and hydrogen cyanide due to photo and collisional*
164 *dissociation and ionizations caused by XUV solar flux and SEP particle flux.*

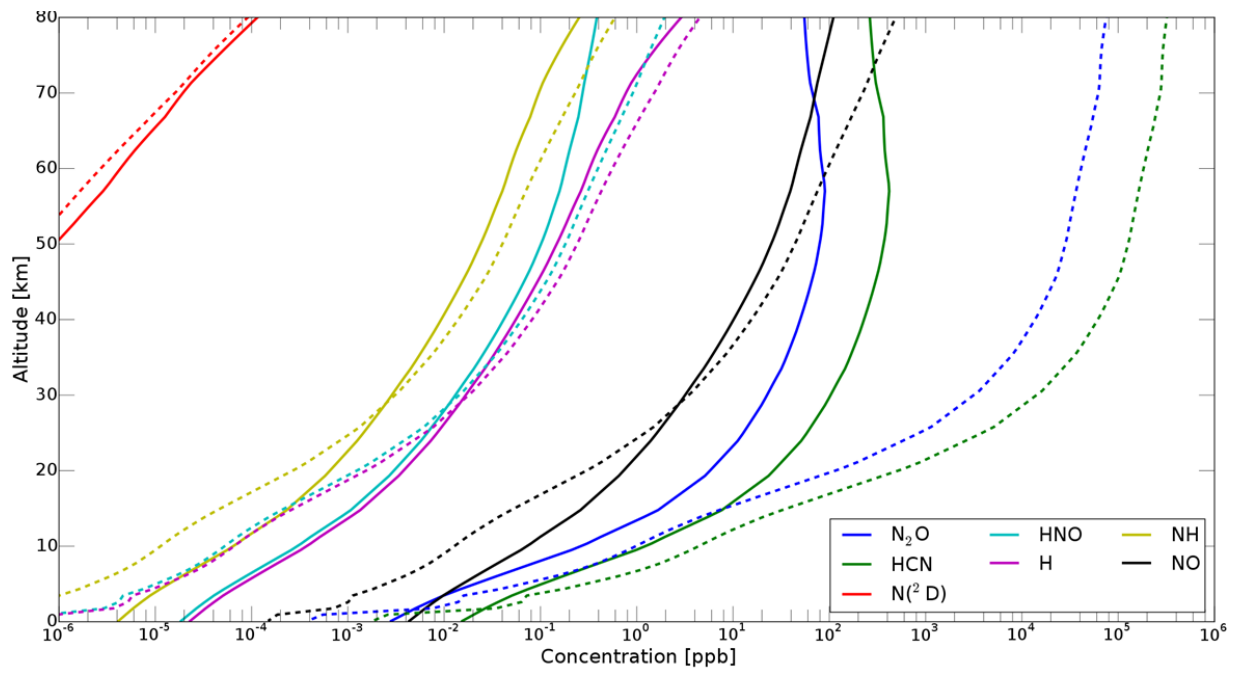
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171 *Figure 4. Radial profiles of the steady-state mixing ratios of various species produced by*
 172 *incoming flux of primary protons and secondary electrons for 10% (dotted lines) and 100%*
 173 *(solid lines) of the maximum photo-destruction rate at 1 PAL (top figure, a) and at 2 PAL*
 174 *(bottom figure, b).*

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261

262 **ACKNOWLEDGEMENTS:** We thank three referees for constructive suggestions that
263 improved the manuscript. This work was supported by NASA GSFC Science Task Group
264 funds. V. Airapetian performed the part of this work while staying at ELSI/Tokyo Tech.

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270 **AUTHOR CONTRIBUTIONS:**

271 V. Airapetian conceived and designed the numerical models, analyzed the data, contributed
272 materials and wrote the manuscript. A. Glocer and G. Gronoff contributed to the

273 development and execution of codes and data analysis. E. Hébrard contributed to the
274 chemistry model and data analysis, W. Danchi contributed to data analysis and proofreading
275 of the paper.

276 **Competing financial interests.**

277 The authors declare no competing financial interests.

278 **FIGURE CAPTIONS.**

279 **Figure 1.** The initial (left panel, a) and the final state (right panel, b) magnetic field lines in
280 white) and the plasma pressure (in color) of the Earth's magnetosphere due to the CME event
281 [14].

282 **Figure 2.** XUV flux of the young Sun at 0.7 Gyr [20].

283 **Figure 3.** The pathway diagram of abiotic production of odd nitrogen and nitrogen-bearing
284 compounds including nitrous oxide and hydrogen cyanide due to photo and collisional
285 dissociation and ionizations caused by XUV solar flux and SEP particle flux.

286 **Figure 4.** Radial profiles of the steady-state mixing ratios of various species produced by
287 incoming flux of primary protons and secondary electrons for 10% (dotted lines) and 100%
288 (solid lines) of the maximum photo-destruction rate at 1 PAL (top figure, 4a) and at 2 PAL
289 (bottom figure, 4b).

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299 **METHODS**

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301 **1. SWMF DESCRIPTION**

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304 In this paper, we utilized the Space Weather Modeling Framework (SWMF) available at

305 Community Coordinated Modeling Center (CCMC) at NASA Goddard Space Flight Center (see

306 at <http://ccmc.gsfc.nasa.gov>). A single-fluid, time dependent fully non-linear 3D

307 magnetohydrodynamic (MHD) code BATS-R-US (Block-Adaptive-Tree Solar-wind Roe-type

308 Upwind Scheme) is a part of SWMF and was developed at the University of Michigan Center of

309 Space Environment Modeling (CSEM). The spine of the SWMF is the BATS-R-US code

310 [30,31], which is coupled to Rice Convection Model (RCM, [31]) to model a propagation and

311 interaction of a model SCME with a magnetosphere and ionosphere of a young Earth. The MHD

312 part of the code calculates the dynamic response of the large-scale magnetospheric plasma to

313 varying solar wind conditions in a self consistent manner by using the block-adaptive wind Roe-

314 type upwind scheme global MHD code [30]. The dynamics of the magnetosphere is described in

315 a Cartesian geometry by using resistive MHD equations. The electromagnetic coupling of the

316 magnetosphere to a conducting ionosphere is handled in a standard way [32]. Specifically, the

317 magnetospheric currents near the inner boundary of the MHD simulation are mapped to the

318 ionosphere where. A potential solver is then used which combines these currents with a

319 conductance map of the ionosphere (including solar and auroral contributions) to produce the

320 electric potential in the ionosphere. That potential is then used to set the electric field and

321 corresponding drift at the magnetospheric simulations inner boundary.

322 The MHD approximation does not provide an adequate description of the inner

323 magnetosphere because energy dependent particle drifts and ring current evolution become

324 important. Here we use the Rice Convection Model, embedded in the MHD simulation, to model
325 this important region [31]. The RCM code is a kinetic plasma model that couples plasma motions
326 in the inner magnetosphere and calculates the energy dependent particle drifts and ring current
327 evolution in the inner magnetosphere. The ring current carries the most of the energy density
328 during magnetic storms and is essential to modeling strong storms. This coupling is crucial for
329 description of solar wind effects on a magnetosphere, because the ionosphere provides closure of
330 magnetospheric currents, which is needed for realistic description of magnetospheric convection
331 and associated electric fields. Thus, we apply a dedicated inner magnetospheric model that is
332 fully coupled to the MHD code for the treatment of the inner magnetosphere. We simulate the
333 magnetospheric cavity (outer and inner magnetosphere) in a computational box defined by the
334 following dimensions $-224R_E < x < 224R_E$, $-128R_E < y < 128R_E$, $-128R_E < z < 128R_E$, where
335 R_E is the radius of the Earth placed at the center of the computational box. The dipole tilt is
336 neglected in this problem. The simulations were carried out using a block adaptive high-
337 resolution grid with the minimum cell size of $1/16 R_E$.

338 The inner boundary is set at $1.25 R_E$ with a density of 100 cm^{-3} . The velocity at
339 the inner body is set to the $\vec{E} \times \vec{B}$ velocity, where \vec{E} is determined from the ionospheric
340 potential and \vec{B} is the Earth's magnetic field. The pressure is set to float. The magnetic field is set
341 in a way that the radial component is the Earth's dipole and the tangential components are
342 allowed to float. The simulation is initialized with a dipole everywhere in the computational
343 domain and a small density, zero velocity, and a finite pressure. The solar wind conditions are set
344 at the upstream boundary and some period of local time stepping is used to get an initial steady
345 state solution. We assume the solar wind input parameters including the three components of

346 interplanetary magnetic field, B_x , B_y and B_z , the plasma density and the wind velocity, V_x ,
 347 using the physical conditions associated with a Carrington- type event as discussed by [33] and
 348 [14], see Figure 1. The time evolution of the plasma pressure (in nPa) and current density (in
 349 microAmps/m²) during the extreme CME event are presented in the attached Movie 1.

350 2. AEROPLANETS MODEL DESCRIPTION

351 We used our sophisticated Aeroplanets model with enhanced chemistry [18] to model the upper
 352 atmospheric region (up to 200 km) in response to young Sun's XUV (X-ray and EUV) emission
 353 from and precipitating electrons and protons due to an SEP event. The model calculates the
 354 photo and collisional (due to protons) dissociation, ionization and photoexcitation processes in
 355 the Earth's atmosphere. The primary photoelectrons are then transported along a magnetic field
 356 line, and the electron impact is computed solving the stationary kinetic Boltzmann equation. This
 357 results in the dissociation, ionization and excitation of the different atmospheric species. The
 358 Aeroplanets code incorporates 117 chemical reactions with the rates presented in Table 1.
 359 To converge to steady state chemical solution for the early Earth atmosphere described in the
 360 Main section was reached after running the code for 6 months of physical time.

361

362 **Table 1. List of Chemical Reactions Used in Our Model**

Reaction	Reaction rate (in cgs units)
$H + CH \rightarrow C + H_2$	$0.124E-09 * (T / 300) ** (0.260E+00)$
$H + CH_2 \rightarrow CH + H_2$	$0.220E-09 * (T / 300) ** (0.320E+00)$

H + e3CH2 -> CH + H2	0.220E-09 * (T / 300) ** (0.320E+00)
H + CH3 -> e3CH2 + H2	0.100E-09 * exp(-0.760E+04 / T)
H + CH4 -> CH3 + H2	0.589E-12 * (T / 300) ** (0.300E+01) * exp(-0.404E+04 / T)
CH + H2 -> e3CH2 + H	0.310E-09 * exp(-0.165E+04 / T)
CH + CH -> C2H + H	2.00E-10
CH + CH -> e3CH2 + C	2.00E-11
CH + e3CH2 -> C2H2 + H	2.00E-10
CH + CH3 -> C2H3 + H	1.00E-11
CH + CH3 -> C2H2 + H + H	1.00E-10
CH + CH4 -> C2H4 + H	0.105E-09 * (T / 300) ** (-0.104E+01) * exp(-0.361E+02 / T)
CH2 + H2 -> CH3 + H	0.880E-10 * (T / 300) ** (0.350E+00)
CH2 + CH4 -> e3CH2 + CH4	0.310E-11 * exp(0.250E+03 / T)
CH2 + CH4 -> CH3 + CH3	0.279E-10 * exp(0.250E+03 / T)
CH2 + C2H2 -> e3CH2 + C2H2	2.30E-10
CH2 + C2H2 -> C3H3 + H	0.760E-10 * (T / 300) ** (-0.300E+00)

$\text{CH}_2 + \text{N}_2 \rightarrow \text{e}^3\text{CH}_2 + \text{N}_2$	$0.110\text{E}-10 * (\text{T} / 300) ** (0.810\text{E}+00)$
$\text{e}^3\text{CH}_2 + \text{H}_2 \rightarrow \text{CH}_3 + \text{H}$	$0.800\text{E}-11 * \exp(-0.450\text{E}+04 / \text{T})$
$\text{e}^3\text{CH}_2 + \text{CH}_4 \rightarrow \text{CH}_3 + \text{CH}_3$	$0.713\text{E}-11 * \exp(-0.505\text{E}+04 / \text{T})$
$\text{CH}_3 + \text{H}_2 \rightarrow \text{CH}_4 + \text{H}$	$0.245\text{E}-13 * (\text{T} / 300) ** (0.288\text{E}+01) * \exp(-0.460\text{E}+04 / \text{T})$
$\text{N}(4\text{S}) + \text{CH} \rightarrow \text{CN} + \text{H}$	$0.140\text{E}-09 * (\text{T} / 300) ** (0.410\text{E}+00)$
$\text{N}(4\text{S}) + \text{e}^3\text{CH}_2 \rightarrow \text{HCN} + \text{H}$	$0.500\text{E}-10 * (\text{T} / 300) ** (0.170\text{E}+00)$
$\text{N}(4\text{S}) + \text{e}^3\text{CH}_2 \rightarrow \text{HNC} + \text{H}$	$0.300\text{E}-10 * (\text{T} / 300) ** (0.170\text{E}+00)$
$\text{N}(4\text{S}) + \text{CH}_3 \rightarrow \text{H}_2\text{CN} + \text{H}$	$5.60\text{E}-11$
$\text{N}(4\text{S}) + \text{CH}_3 \rightarrow \text{HCN} + \text{H} + \text{H}$	$6.00\text{E}-12$
$\text{N}(4\text{S}) + \text{NH} \rightarrow \text{N}_2 + \text{H}$	$0.250\text{E}-10 * (\text{T} / 300) ** (0.170\text{E}+00)$
$\text{N}(4\text{S}) + \text{CN} \rightarrow \text{C} + \text{N}_2$	$0.900\text{E}-10 * (\text{T} / 300) ** (0.420\text{E}+00)$
$\text{N}(4\text{S}) + \text{H}_2\text{CN} \rightarrow \text{N}_2 + \text{e}^3\text{CH}_2$	$4.00\text{E}-11$
$\text{N}(4\text{S}) + \text{H}_2\text{CN} \rightarrow \text{HCN} + \text{NH}$	$5.00\text{E}-12$
$\text{N}(2\text{D}) \rightarrow \text{N}(4\text{S})$	$2.30\text{E}-05$
$\text{N}(2\text{D}) + \text{H}_2 \rightarrow \text{NH} + \text{H}$	$0.420\text{E}-10 * \exp(-0.880\text{E}+03 / \text{T})$
$\text{N}(2\text{D}) + \text{CH}_4 \rightarrow \text{NH} + \text{CH}_3$	$0.130\text{E}-10 * \exp(-0.755\text{E}+03 / \text{T})$

N(2D) + CH4 -> CH2NH + H	0.350E-10 * exp(-0.755E+03 / T)
N(2D) + N2 -> N(4S) + N2	0.100E-12 * exp(-0.520E+03 / T)
N(2D) + NH3 -> N2H2 + H	5.00E-11
N(2D) + HCN -> CH + N2	5.00E-11
N(2D) + HNC -> CN2 + H	2.00E-11
N(2D) + HNC -> CH + N2	2.00E-11
NH + H -> N(4S) + H2	0.220E-11 * (T / 300) ** (0.155E+01) * exp(-0.103E+03 / T)
NH + CH -> HCN + H	5.00E-11
NH + CH -> HNC + H	5.00E-11
NH + e3CH2 -> H2CN + H	3.00E-11
NH + e3CH2 -> HCN + H + H	3.00E-11
NH + e3CH2 -> HNC + H2	5.00E-12
NH + CH3 -> CH2NH + H	0.130E-09 * (T / 300) ** (0.170E+00)
NH + NH -> N2 + H + H	2.00E-10
NH + NH2 -> N2H2 + H	0.100E-09 * (T / 300) ** (0.170E+00)
NH + CN -> CN2 + H	1.00E-10

NH + CN -> N2 + CH	1.00E-10
NH2 + H2 -> NH3 + H	0.209E-11 * exp(-0.428E+04 / T)
NH2 + H -> NH + H2	0.200E-10 * exp(-0.240E+04 / T)
NH2 + CH4 -> NH3 + CH3	0.399E-13 * (T / 300) ** (0.359E+01) * exp(-0.454E+04 / T)
NH2 + C2H2 -> NH3 + C2H	0.111E-12 * exp(-0.185E+04 / T)
NH2 + C2H3 -> NH3 + C2H2	2.00E-11
NH2 + C2H3 -> SOOTN + H	8.00E-11
NH2 + H2CN -> HCN + NH3	0.540E-10 * (T / 300) ** (-0.110E+01) * exp(-0.600E+02 / T)
NH3 + H -> NH2 + H2	0.423E-13 * (T / 300) ** (0.393E+01) * exp(-0.406E+04 / T)
NH3 + CH -> CH2NH + H	0.169E-09 * (T / 300) ** (-0.560E+00) * exp(-0.280E+02 / T)
NH3 + CH3 -> NH2 + CH4	0.510E-13 * (T / 300) ** (0.286E+01) * exp(-0.734E+04 / T)
CN + H2 -> HCN + H	0.412E-12 * (T / 300) ** (0.287E+01) * exp(-0.820E+03 / T)
CN + CH -> HCN + C	0.100E-09 * (T / 300) ** (-0.170E+00)
CN + e3CH2 -> HCN + CH	5.00E-11

CN + e3CH2 -> CHCN + H	5.00E-11
CN + e3CH2 -> C2N + H2	5.00E-11
CN + CH3 -> CH2CN + H	1.00E-10
CN + CH4 -> HCN + CH3	0.620E-11 * exp(-0.721E+03 / T)
CN + NH3 -> HCN + NH2	0.277E-10 * (T / 300) ** (-0.114E+01)
CN + HCN -> C2N2 + H	0.430E-12 * (T / 300) ** (0.171E+01) * exp(-0.770E+03 / T)
CN + HNC -> C2N2 + H	2.00E-10
HCN + CH -> CHCN + H	0.140E-09 * (T / 300) ** (-0.170E+00) * exp(-0.0 / T)
HCN + CH -> C2N + H2	0.140E-09 * (T / 300) ** (-0.170E+00) * exp(-0.0 / T)
HCN + 1C2 -> C3N + H	0.200E-09 * (T / 300) ** (0.170E+00) * exp(-0.0 / T)
HNC + H -> HCN + H	0.300E-10 * exp(-0.800E+03 / T)
H2CN + H -> HCN + H2	6.00E-11
H2CN + CH3 -> CH4 + HCN	3.00E-11
O(3P) + CH2 -> HCO + H	1.00E-11
O(3P) + CH2 -> CO + H + H	5.00E-11
O(3P) + CH2 -> CO + H2	6.00E-11

$O(3P) + e3CH2 \rightarrow HCO + H$	1.00E-11
$O(3P) + e3CH2 \rightarrow CO + H + H$	5.00E-11
$O(3P) + e3CH2 \rightarrow CO + H2$	6.00E-11
$O(3P) + CH3 \rightarrow CO + H2 + H$	2.90E-11
$O(3P) + CH3 \rightarrow H2CO + H$	1.10E-10
$O(3P) + NH \rightarrow NO + H$	6.60E-11
$O(3P) + HNO \rightarrow OH + NO$	3.80E-11
$O(1D) + H2 \rightarrow OH + H$	1.10E-10
$O(1D) + CH4 \rightarrow OH + CH3$	1.05E-10
$O(1D) + CH4 \rightarrow CH3O + H$	3.50E-11
$O(1D) + CH4 \rightarrow H2CO + H2$	7.50E-12
$O(1D) + N2 \rightarrow O(3P) + N2$	2.15E-11
$OH + H2 \rightarrow H2O + H$	$0.280E-11 * \exp(-0.180E+04 / T)$
$OH + CH4 \rightarrow H2O + CH3$	$0.185E-11 * \exp(-0.169E+04 / T)$
$OH + N(4S) \rightarrow NO + H$	4.50E-11
$OH + N(2D) \rightarrow NO + H$	4.50E-11

OH + CO -> CO2 + H	1.30E-13
H2O + CH -> H2CO + H	$0.280E-10 * (T / 300) ** (-0.122E+01) * \exp(-0.120E+02 / T)$
H2O + N(2D) -> OH + NH	4.50E-11
H2O + N(2D) -> HNO + H	5.00E-12
CO2 + N(2D) -> CO + NO	$0.100E-10 * \exp(-0.100E+04 / T)$
NO + CH -> HCN + O(3P)	$0.100E-09 * (T / 300) ** (-0.130E+00)$
NO + CH -> NCO + H	$0.300E-10 * (T / 300) ** (-0.130E+00)$
NO + CH -> CO + NH	$0.300E-10 * (T / 300) ** (-0.130E+00)$
NO + CH -> OH + CN	$0.100E-10 * (T / 300) ** (-0.130E+00)$
NO + e3CH2 -> HNCO + H	$0.210E-11 * \exp(0.554E+03 / T)$
NO + e3CH2 -> CO + NH2	$0.300E-12 * \exp(0.554E+03 / T)$
NO + N(4S) -> O(3P) + N2	$0.400E-10 * (T / 300)$
NO + N(2D) -> O(3P) + N2	$0.600E-10 * (T / 300)$
NO + NH -> N2O + H	$0.290E-10 * (T / 300) ** (-0.300E+00) * \exp(0.770E+02 / T)$
NO + NH -> OH + N2	$0.120E-10 * (T / 300) ** (-0.300E+00) * \exp(0.770E+02 / T)$

HNO + H -> NO + H2	$0.310E-10 * \exp(-0.500E+03 / T)$
HNO + N(2D) -> NO + NH	5.00E-11
N2O + CH -> NO + HCN	$0.150E-10 * \exp(0.257E+03 / T)$
N2O + N(2D) -> N2 + NO	$0.150E-10 * \exp(-0.570E+03 / T)$
H + H -> H2	$(0.914E-32 * (T / 300) ** (-0.600E+00) * \exp(-0.0 / T)) * [M] / (1 + (0.914E-32 * (T / 300) ** (-0.600E+00) * \exp(-0.0 / T)) * [M] / (0.100E-09 * \exp(-0.0 / T)))$
H + CH3 -> CH4	$(0.890E-28 * (T / 300) ** (-0.180E+01) * \exp(-31.8 / T)) * [M] / (1 + (0.890E-28 * (T / 300) ** (-0.180E+01) * \exp(-31.8 / T)) * [M] / (0.320E-09 * (T / 300) ** (0.133E+00) * \exp(-2.54 / T)))$
HCN + H -> H2CN	$(0.100E-33 * \exp(-0.0 / T)) * [M] / (1 + (0.100E-33 * \exp(-0.0 / T)) * [M] / (0.980E-11 * \exp(-2080.0 / T)))$

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377 **Code availability.** BATS-R-US code as a part of SWMF used to generate the scenario shown in

378 Figure 1 can be accessed through GSFC's CCMC run-on-request web site at

379 <http://ccmc.gsfc.nasa.gov>. We have opted not to make the computer code, Aeroplanets,

380 associated with this paper available because we are currently in the phase of adopting it for free

381 access at Exoplanetary part of CCMC's web site. The authors declare that model data supporting

382 the findings of this study are available within the article and its supplementary information.

383 Other model related data will be available on request from the authors.