| 1 | Soil enzyme activity and stoichiometry along a gradient of vegetation restoration at the |
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| 2 | Karst Critical Zone Observatory in southwest China |
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17 Abstract

18 The 'Grain for Green Programme (GGP)' was implemented in the 1990s as a solution to the 19 extreme degradation of karst landscapes that cover one-third of China, largely caused by 20 decades of poorly managed intensive agriculture. The recovery of soil functions is key to the 21 success of ecosystem regeneration of abandoned croplands where the carbon (C) and nutrient 22 cycles have been severely perturbed by cultivation. However, an ecological 'tipping point' 23 beyond which soil functions are unrecoverable in manageable timescales may have been passed 24 in the fragile, subtropical karst ecosystem. The aim of this study was to use the activity of key 25 enzymes for C, nitrogen (N) and phosphorus (P) acquisition in the soil as a proxy for the 26 biological response to vegetation restoration after agricultural abandonment in a severely 27 degraded karst catchment at the Karst Critical Zone Observatory in Guizhou Province. In 2016, 28 a space-for-time approach was used to establish a chronosequence of vegetation recovery: 29 sloping cropland < recently abandoned sloping cropland < shrubland < secondary (regenerated) 30 forest < primary (natural) forest. Soils were sampled from the surface to the bedrock (up to 80 31 cm depth) in each recovery phase. The activity of all enzymes in the top 0-30 cm depth 32 increased after abandonment and was positively correlated with soil nutrient and water contents. 33 Nitrogen deficiencies were indicated by the reduced ratios of C- relative to N-hydrolase activity 34 and the increased ratios of N- relative to P-hydrolase activity in the abandoned croplands and 35 shrublands. Phosphorus deficiencies were indicated by the reduced ratios of N- relative to P-36 hydrolase activity and C- relative to P-hydrolase activity in the soils of the shrubland and 37 secondary forest compared to the primary forest. Our results revealed that near-to-natural 38 biological soil function was recoverable as vegetation naturally restored and suggested that the 39 rate of recovery may be accelerated by managed nutrient amendments during the early stages 40 after abandonment. This new information may help to inform the managed regeneration of 41 degraded agricultural land in nutrient-poor, sub-tropical environments.

42 Keywords: soil organic carbon, nitrogen, phosphorus, enzyme activity, vegetation restoration;
43 karst ecosystem

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45 1 INTRODUCTION

46 Soil degradation caused by human perturbation causes unsustainable losses in net primary 47 productivity (NPP) that threatens human wellbeing and development across the world (Bai et 48 al., 2011). The most recent Global Assessment of Land Degradation and Improvement 49 concluded that land degradation is most conspicuous in the rapidly developing, humid south of 50 China, by comparison with the dry lands of the north and west (Bai et al., 2009). Over the last 51 50 years, unsustainable intensification of agriculture, exploitation of mineral resources and 52 encroaching urbanization have caused progressive and severe degradation in substantial areas 53 of the fragile karst ecosystems of southwest China, one of the largest karst regions in the world 54 (Montgomery, 2007). The characteristic steeply sloping topographies of karst landscapes have 55 shallow soil profiles that are vulnerable to erosion and leaching (Delang & Yuan, 2015; Quine 56 et al., 2017). In many areas of the karst landscapes of southwest China, a 'tipping point' appears 57 to have been passed as basement rock is exposed and 'rocky desertification' dominates because 58 soil is being lost as a consequence of accelerated rates of erosion caused by human perturbation 59 (Wang et al., 2004). Thus, there is a pressing need to better understand and quantify the changes 60 that occur in soil quality and ecosystem function before this extreme state is reached so that 61 appropriate management stratagies, if any, can be informed (Quine et al., 2017).

62 In recognition of the severity of the degradation problem in the karst ecosystems of Southwest 63 China, the Chinese Government implemented the 'Grain for Green Programme' (GGP) in the 64 region in the late 1990s. Croplands on slopes greater than 25° were abandoned and vegetative 65 succession allowed to proceed naturally from non-woody plants to woody shrubs and ultimately 66 to mature secondary forest, with the ultimate objective of ecosystem restoration to conditions 67 similar to pristine primary forest (Xu et al., 2006). The regeneration of plant communities under 68 GGP in degraded karst landscapes has been largely successful, indicated by increased net 69 primary productivity (NPP) and soil organic matter (SOM) stocks, and the stabilization of soil 70 profiles by vegetative canopy cover and roots that moderate further loss by erosion (Chang et 71 al., 2017; Ren et al., 2017; Tong et al., 2018). Soil organic matter contains soil organic carbon 72 (C) and nutrients (including nitrogen (N) and phosphorus (P)) and is recognized as a key 73 indicator of soil health and productivity in agricultural systems (Dungait et al., 2012a). The turnover of SOM and release of nutrients is mediated by the soil microbial community (e.g. bacteria and fungi including mycorrhiza) that also play a role in soil stabilization through the formation of soil aggregates, through the process of 'soil self-organisation' (Young & Crawford, 2004), increasing the potential for successful vegetation colonization and succession by improving soil quality, i.e. nutrient and water retention and supply and diffusion of gases through soil pores (Brahma et al., 2018).

80 Ecosystem responses to land use change are controlled by the potential to adapt to changing C 81 and nutrient supply, described as 'stoichiometric flexibility' by Sistla & Schimel (2012). 82 Resource availability controls biological processes that subsequently effect further change in 83 supply and acquisition to meet the basic metabolic requirements of all organisms (Elser et al., 84 1996). Cleveland & Liptzin (2007) recognized constrained atomic C:N:P ratios in soils and the 85 soil microbial biomass at the global scale, that nevertheless showed significant variation in C:N, 86 C:P and N:P ratios between vegetation types, e.g. forest versus grassland. This approach has 87 been applied to investigate soil microbial responses to land use change in Chinese karst 88 ecosystems. For example, Li et al. (2012) determined significant variation in the C:N:P ratios 89 of soil and soil microbial biomass in typical karst mountain, low hill and lowland landscapes 90 that had been caused by human perturbation. Hu et al. (2016) described coincident increases in 91 soil and microbial C:N ratios in reforested karst soils that demonstrated an adaptive response 92 by the soil microbial community to changes in substrate resource stoichiometry. Soils in natural 93 karst ecosystems are characteristically nutrient-limited because of underlying geology, and 94 many studies have explored nutrient limitations using N:P ratios of soil (Jiao et al., 2013), plant 95 foliage (Zhang et al., 2015a) and microbial biomass (Tischer et al., 2014). Typical changes in 96 soil N:P ratios caused by a shift from N- to P-limitation during vegetation restoration are 97 observed in karst ecosystems (Zhang et al., 2015a) as N leaches readily in the early stages after 98 abandonment with sparse vegetation coverage (Song et al., 2017), and P adsorbs strongly to 99 calcium (Ca) weathered from the bedrock (Vitousek et al., 2010).

100 Most studies of soil processes focus on surface soil horizons where soil microorganisms are 101 more abundant in response to greater inputs of organic matter from decomposing surface and 102 root litter and rhizoexudates (Gocke et al., 2017; Peng & Wang, 2016). Soil enzyme activities

103 decline exponentially with depth in the soil profiles of temperate grasslands and tropical forests 104 (Blume et al., 2002; Stone et al., 2014). However, belowground competition for nutrients and 105 water by plant roots is recognized in many ecosystems, and symbiotic mycorrhizae scavenge 106 throughout the soil in resource-limited soils, e.g. in arable soils after long term cultivation and 107 litter removal (Kautz et al., 2013). In karst systems, the depth of soils is generally considered 108 to be very shallow due to slow rates of pedogenesis from highly soluble bedrock, but the 109 potential for soil accretion in deep geological formations in the epikarst is substantial but poorly 110 understood. The combination of the unique geology of karst ecosystems and the limited 111 understanding of soil enzyme activity and their ratios in karstic subsoils suggests a unique 112 opportunity to understand the role of subsoil nutrient cycling and supply during vegetation 113 restoration in these fragile ecosystems.

114 Soil microorganisms and some plant roots produce extracellular enzymes that catalyze the 115 mineralization of SOM, releasing nutrients into soil solution that become available for take-up 116 by plant roots (Burns et al., 2013). The resource allocation strategies of soil microorganisms to 117 metabolic processes (including enzyme biosynthesis) respond to the availability of substrates 118 and are constrained by their stoichiometric requirements to maintain metabolic and nutrient 119 balances (Sinsabaugh et al., 2002). Therefore, soil enzyme activity has been used as a sensitive 120 indicator of biological function for several decades (e.g. Dick, 1994) and extracellular enzyme 121 activities are commonly interpreted as indicators of microbial nutrient demand (Sinsabaugh et 122 al., 2008). Soil hydrolases preferentially catalyze nutrient-rich 'labile' SOM related directly to 123 meeting the C, N and P nutrient requirements of soil microorganisms and plants, whereas soil 124 oxidase activities are related to the turnover of poor quality 'recalcitrant' polyphenolic organic 125 matter, e.g. lignin, terpenoids and humified organic matter (Sinsabaugh & Shah, 2011). Thus, when investigating proximity to a 'tipping point' in ecosystem function, or ecosystem 126 127 regeneration away from a degraded state, the activity of ubiquitous enzymes for C and nutrient 128 cycling in soils should provide a reliable indicator of change in soil biological function. A global 129 scale meta-analysis of 40 terrestrial ecosystems revealed that the ratios of soil lnC: lnN: lnP 130 nutrient acquisition hydrolase activities were close to 1:1:1, indicating the microbial nutrient 131 acquisition capacities were generally similar (Sinsabaugh et al., 2008). Xu et al. (2017)

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132 investigated C, N and P nutrient acquisition hydrolase activities in forests along the North-133 South Transect in eastern China (NSTEC) using the ratios of lnC:lnN:lnP. They concluded that 134 the reduced ratios of lnC:lnP and lnN:lnP tropical and subtropical forest soils indicated 135 predominant P-limitation compared with temperate forests and the global average. Thus, 136 changes in extracellular enzyme activity and their ratios may be used as a tool to explore the 137 constraints on microbial biomass (Sinsabaugh et al. 2008) by providing a functional assessment 138 of nutrient availability and limitations for vegetation restoration due to inadequate soil fertility 139 (Waring et al., 2014).

140 The overall objective of this study was to qualify and quantify how the activities and ratios of 141 soil enzymes through soil profiles responded to changes in nutrient availability during 142 vegetative restoration of a degraded watershed under GPP compared to a non-degraded pristine 143 forest in the karst ecosystem of Southwest China. We hypothesized that N would be limiting in 144 the early stages of restoration when the vegetation cover was sparse because of low N inputs 145 and strong nitrate leaching, and soil C:N acquiring hydrolase activities would also be reduced 146 in the early stages compared to later stages of restoration due to larger organic matter input 147 from increasing NPP. We also postulated that soil N:P acquiring hydrolase activities would be 148 poor because karst soils are naturally P-limited due to low inputs of P from parent rock 149 weathering and strong adsorption by Ca.

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151 2 MATERIALS AND METHODS

152 **2.1** Study sites

Two field sites in close proximity with the same soil type and climate regime (Chenqi watershed and Tianlong Mountain) were selected for study at the Karst Critical Zone Observatory in Puding County, Guizhou Province, Southwest China, encompassing five recovery phases along a chronosequence of recovery in the karst ecosystem. The sloping cropland, recently abandoned sloping cropland, shrubland and secondary (regenerated) forest were in the Chenqi watershed, and the primary (pristine) forest was on Tianlong Mountain. The two field sites are representative of the regional karst ecosystem and subject to a subtropical monsoon climate. The annual mean precipitation was 1315 mm, with an annual mean temperature of ~15.1 °C (Liu et al., 2016). The soil in both areas was dominated by Mollic Inceptisols that originated from the limestone bedrock of the Middle Triassic Guanling Formation (Lu et al., 2014). Table 1 provides the landscape and management characteristics of the sites and the dominant vegetation species on each land use.

The Chengi catchment (26°15'37"- 26°15'40" N, 105°46'11"- 105°46'29" E) ranges in 165 166 elevation from 1100 and 1600 m above sea level (asl), with an area of 1.29 km². Intensive 167 cultivation (predominantly maize, oil seed rape and soybean in rotation with regular application 168 of urea as a single compound or combined with manure) had been practised on the hillslopes 169 since the 1960s. The GGP was introduced here in the 1990s to encourage abandonment of large 170 areas of severely degraded sloping cropland on middle and upper slopes allowing regeneration 171 to woody shrubland (2-3 years) and then secondary forest (10 years +). Tianlong Mountain (26° 172 $14'44'' - 26^{\circ}14'48''N$, $105^{\circ}45'40'' - 105^{\circ}45'46''E$) is predominately an undisturbed, pristine 173 primary forest with no history of farming, with an elevation between 1421 and 1503 m asl.

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175 2.2 Soil sampling

176 In July 2016, we used a space-for-time approach to establish a chronosequence of change 177 following agricultural land abandonment under GGP. Soils were sampled from four replicate 178 $10 \text{ m} \times 20 \text{ m}$ plots in each of the five recovery phases: sloping cropland < recently abandoned 179 sloping cropland < shrubland < secondary (regenerated) forest < primary (natural) forest in the 180 Chenqi catchment and primary forest on Tianlong Mountain (i.e. 20 plots in total). A soil auger 181 with a diameter of 2 cm was used to collect soil samples from different soil depths intervals (i.e. 182 0-10, 10-30, 30-50, 50 cm to the bedrock). Soil depth did not extend below 50 cm in the 183 abandoned croplands and shrublands. To account for spatial heterogeneity, ten soil cores were 184 randomly collected within each plot and composited into a single soil sample. Each composite 185 sample was stored at 4 °C until analysis for soil water content (SWC), pH, and concentrations 186 of dissolved organic carbon (DOC), available nitrogen (AN), available phosphorus (available P) and soil enzyme activities. Subsamples were air dried, homogenized and ground to 0.15 mm
for soil organic carbon (SOC), total nitrogen (TN), and total phosphorus (TP) determination.

189 2.3 Soil environmental parameters

190 Percentage SWC was measured by oven drying fresh, sieved (2mm) soil at 105 °C to constant 191 weight (Bao, 2008). Soil pH was measured in a 1:2.5 v:v soil-water suspension using a glass electrode (Bao, 2008). SOC (g kg⁻¹) was determined by dichromate oxidation and titration with 192 193 ferrous ammonium sulfate (Bao, 2008). TN (g kg⁻¹) was determined by dry combustion using 194 an element analyzer (Elementar, Vario Max CN, Germany) (Bao, 2008). DOC (mg kg⁻¹) was 195 extracted in distilled water at a soil: distilled water ratio of 1:5 and quantified using an organic 196 carbon analyzer (Elementar, Liqui TOC II, Germany) (Bao, 2008). TP (mg kg⁻¹) was measured 197 spectrophotometrically with a continuous flow auto-analyzer (Bran Luebbe, AA3, Germany) 198 after digestion with H₂SO₄-HClO₄ (Bao, 2008). AN (mg kg⁻¹) was determined after extraction with 1 mol L^{-1} KCl, and available P after extraction with 0.5 mol L^{-1} NaHCO₃, by continuous 199 200 flow analyser (Bran Lubbe, AA3, Germany) (Bao, 2008).

201 2.4 Enzyme assays

202 The potential activities of three C-acquiring enzymes (β -1,4-glucosidase, β G, EC 3.2.1.21; β -203 D-1,4-cellobiosidase, CBH, EC 3.2.1.91; β-1,4-xylosidase, βX, EC 3.2.1.37), two N-acquiring 204 enzymes (β-N-acetyl glucosaminidase, NAG, EC 3.1.6.1; leucine aminopeptidase, LAP, EC 205 3.4.11.1) and one P-acquiring enzyme (alkaline phosphatase, AP, EC 3.1.3.2) in the soil 206 samples were measured following a modified fluorescence method (German et al., 2011; Trap 207 et al., 2012). In brief, 1 g of fresh soil was homogenized in 125 ml 2-morpholin-4-208 ylethanesulfonic acid (MES) buffer (pH = 6.5) or 2-Amino-2-(hydroxymethyl) propane-1,3-209 diol (Tris) buffer (pH = 7.5). The pH of the buffers was selected because they were similar to 210 the soil pH, according to German et al. (2011). 200 µl of homogenate and 50 µl of substrate 211 was added to a 96-well black microplate. Eight replicates for each soil sample, blanks, negative 212 controls and quench standards were analyzed. The microplates were incubated in the dark at 213 25 °C for 4 hours according to Saiya-Cork et al. (2002). After incubation, 10 µl of 1 mol L⁻¹ 214 NaOH was added to each well to terminate the reactions, and fluorescence values were 215 measured using a microplate fluorometer (Synergy^{H4}, BioTek, USA) with excitation and 216 emission filters of 365 and 450 nm, respectively. The absolute hydrolase activities were 217 expressed in units of nmol g^{-1} soil h^{-1} after correcting for negative controls and quenching. The 218 ratios of C, N and P nutrient acquiring hydrolase activities were compared using 219 lnBG:ln(NAG+LAP):lnAP.

220 Soil oxidase activities (i.e. Peroxidase, POD, EC 1.10.3.2; Polyphenol oxidase, PPO, EC 221 1.11.1.7) were assayed colourimetrically using L-3,4-dihydroxyphenylalanine (L-DOPA) 222 following a modified microplate protocol optimized for in situ pH conditions (German et al., 223 2011; Saiya-Cork et al., 2002). In brief, 600 μ l of homogenate and 150 μ l of substrate were 224 added to deep 96 well microplates. For measuring the POD activities, an extra 10 μ l of 0.3% 225 H₂O₂ was added, then incubated for 5 hours at 25°C in the dark. After incubation, the 226 microplates were centrifuged at 3000 rpm for 3 minutes, and 250 µl of liquid supernatant was 227 transferred to a 96 well transparent microplate. The absorbance values were measured at 460 228 nm by microplate spectrophotometer (Synergy^{H4}, BioTek, USA). The absolute oxidase activities were expressed in units of μ mol g⁻¹ soil h⁻¹. The substrates and functions of the 229 230 enzymes are summarized in Table S1.

231 2.5 Statistical analyses

232 All results are reported as means ± standard errors, and analysis was completed using SPSS 21 233 software (IBM SPSS Statistics, USA). All variables of data were tested for normal distribution 234 and homogeneity of variance, with no transformations necessary. Two-way ANOVA was used 235 to identify the significant differences in the activities and stoichiometry of C, N, and P nutrient-236 acquiring enzymes and the soil properties of different soil depths among land use types. 237 Duncan's test was used for Post-Hoc comparisons to identify the significant sub-sets, if 238 apparent. The relationships between the soil properties and the soil enzyme ratios were assessed 239 using Pearson correlations in SPSS. Results were deemed statistically significant when p < 0.05. 240 The figures were plotted in OriginPro 2017 (Originlab Corporation, USA).

241

242 3 RESULTS

243 **3.1** Soil properties

Soil moisture, SOC and TN contents were significantly affected by land use types, soil depth 244 245 and their interactions (Table 2, S2, S3, S4, p < 0.05). The SWC (%) was larger in the top 30 cm 246 of soil from the secondary (+61%, p < 0.05) and primary forests (+50%, p < 0.05) compared to 247 the sloping croplands. The SOC (+119%, p < 0.05) and TN contents (+70%, p < 0.05) were 248 greater in the secondary forests than in the sloping croplands, and ~184% and 150% larger in 249 the primary forests compared to the sloping croplands respectively (p < 0.05). The DOC 250 contents were increased in the 0-10 cm soil depth from the shrublands (+181%, Table 2, S3, p251 < 0.05) and secondary (+73%, p < 0.05) and primary forests (+185%, p < 0.05) relative to the 252 abandoned croplands. The AN content was increased in the 0-10 cm soil depth from the 253 shrublands (+104%, p < 0.05) and secondary (+141%, p < 0.05) and primary forests (+193%,

254 p < 0.05) compared to the abandoned croplands.

TP contents were significantly decreased in the shrublands and secondary forests compared to the other vegetation types, and available P contents were significantly larger in the primary forests than in the other vegetation types (Table S3, p < 0.05). The SOC:TN, SOC:TP and TN:TP ratios were increased in the shrublands and secondary forests relative to the other land use types throughout the soil profile (Table S3, p < 0.05).

- Along the soil profiles, total nutrient contents (i.e. SOC, TN, and TP) declined with depth (Table
- 261 S4, p < 0.05). The AN contents did not vary along the soil profiles in the abandoned croplands
- and shrublands (Table 2, p < 0.05). The SWC was greater in the top 10 cm compared to below
- 263 10 cm in the secondary and primary forests (p < 0.05).
- 264 **3.2** Soil enzyme activity and their ratios
- The results of two-way ANOVA indicated that soil depth, land use type and their interactions could all influence soil hydrolase and oxidase activities (Fig. 1a-f, Table S2, p < 0.05). The soil hydrolase activities were affected by the land use type at depths of 0-30 cm. Moreover, the BG,
- 268 BX and CBH activities were approximately 3.6, 2.0 and 2.3 times increased, respectively, in
- the primary forest compared to the abandoned cropland (Fig. 1a, b, c, p < 0.05). The NAG, AP
- and LAP activities in the primary forests were 11.9, 4.1 and 1.0 times larger, respectively, than

in the abandoned croplands (Fig. 1d, e, f, p < 0.05). The soil PPO and POD activities in the 0-30 cm soil layers were greater in the secondary and primary forests relative to the other land use types (Fig. 1g, h, p < 0.05).

274 The results of two-way ANOVA indicated that soil depth and land use type influenced soil 275 enzyme activity ratios (Fig. 2, 3). The average ratio of soil lnBG:ln(NAG+LAP):lnAP among 276 the land use types at the four depth intervals was 1:1.2:1.4. The ratios of soil 277 InBG:In(NAG+LAP) were less in the abandoned croplands (0.73) and shrublands (0.74) 278 compared to the other land uses (0.82~0.84, Fig. 2a, p < 0.05). The ratios of 279 ln(NAG+LAP):lnAP were increased in the abandoned croplands (0.84) compared to the other 280 land uses (0.70~0.78; Fig. 2b, p < 0.05). The ratios of lnBG:lnAP were reduced in the 281 shrublands (0.55) relative to the other land use types (0.60~0.64; Fig. 2c, p < 0.05).

The soil hydrolase and oxidase (PPO and POD) activities decreased in the top 30 cm of soil but did not differ below 30 cm (Fig. 1a-h, Table S4). Soil NAG, LAP, and AP activities decreased with soil depth in the secondary and primary forests, but not in the other land uses types (Fig. 1d, e, f, p < 0.05). Soil lnBG:lnAP ratios were greater in the top 30 cm than below 30 cm (Fig. 3, p < 0.05).

287 **3.3** Correlations between the soil enzyme activities and soil properties

The results of Pearson correlation analysis showed that soil hydrolase and oxidase activities, and their stoichiometric relationships (i.e. lnBG:ln(NAP+LAP) and lnBG:lnAP) were generally positively related to SWC, SOC, DOC, TN, AN and TP. In addition, soil hydrolase activities were positively related to available P, SOC:TN and SOC:TP. The ratio of lnBG:ln(NAP+LAP) was negatively correlated with pH, while lnBG:lnAP was positively correlated to SOC:TN(Table 3, p<0.05).

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295 4 DISCUSSION

4.1 Soil enzyme activity in soil profiles at different stages of vegetation recovery

297 The synthesis of enzymes by microorganisms in response to environmental stress associated

298 with land management change is likely to be rapid (Burns et al., 2013). In some instances, 299 changes in microbial populations or activity can precede detectable changes in soil physical and 300 chemical properties, thereby providing an early sign of soil improvement or an early warning 301 of soil degradation (Pankhurst et al. 1995). In this study, the differences between vegetation 302 recovery stages (i.e. sloping farmland, abandoned farmland, shrubland, secondary forest and 303 primary forest) were clear across all of the soil enzymes considered in this study (see Fig. 1). 304 Soil C, N, and P nutrient-acquisition enzyme activities were all significantly positively 305 correlated with total soil nutrients (i.e. SOC, TN, and TP) and available nutrient content (i.e. 306 DOC, AN and available P) as expected in nutrient-limited karst ecosystems (Xu et al., 2015; 307 Ren et al. 2016). The relationship between vegetation restoration and associated increases in 308 NPP and SOM content are widely recognised (e.g. Liu et al., 2015) because there is a direct 309 enhancement of the supply of substrate to soil microorganisms, and an indirect increase in 310 supply via elevated soil hydrolase activity (i.e. positive feedback). As expected, and consistent 311 with previous studies (e.g. Peng & Wang, 2016; Stone et al., 2014), land use type had the 312 predominant effect on the soil enzyme activities in the top 30 cm of the soil profile where input 313 from plant litter and roots is greatest and nutrients are most abundant (Jackson et al., 1996; Lee 314 et al., 2014), compared to below 30 cm depth where SOM content is reduced and of poorer 315 quality and oxygen supply may be limiting (Fontaine et al., 2007; Schrumpf et al., 2013; Bai et 316 al., 2015).

317 A positive relationship between in situ SWC and soil enzyme activity observed herein (Table 318 3) supports the widely reported importance of soil moisture for biological function, including 319 soils in the Chinese karst ecosystem (Chen et al. 2017). It is likely that this effect was a product 320 of enhanced water holding capacity associated with increasing organic matter inputs from 321 greater litter inputs above- and below-ground (roots and rhizoexudation) as NPP increased 322 along the recovery gradient from arable to secondary forest, combined with the enhanced 323 abundance and activity of soil microorganisms and subsequent improvements in soil 324 aggregation (Zhu et al., 2012; Wang et al, 2015). Soil aggregation improves soil pore 325 connectivity through an increase in soil macropores, providing pathways for water, nutrients 326 and air that supports biological function, including root proliferation. Enzyme biosynthesis and

secretion is subsequently enhanced as the soil microbial community proliferates under more
favourable physical and chemical conditions (Keeler et al., 2009). We recognize that plant roots
in the study plots may also secrete enzymes directly into the soil to acquire nutrients (Dakora
& Phillips, 2002), but we did not differentiate between sources in this study.

331 In the sloping croplands, soil hydrolase and oxidase activities in the 0-30 cm depth were 332 generally relatively increased in the actively cultivated soils prior to abandonment under GPP. 333 We suggest that the regular, largely manual tillage (by hoe), fertilizer application and crop 334 residues increased the porosity of the top soil horizon directly, leading to an enhancement of 335 aerobic enzyme activity (Bandick & Dick, 1999; Munkholm et al., 2016). By contrast, the 336 recently abandoned croplands may have experienced a rapid diminishment in nutrient 337 availability, fast decomposition of crop residues and SOM (i.e. priming; Kuzyakov et al., 2000), 338 and soil compaction caused by slaking (i.e. collapse of soil structure by water inundation) in 339 the recently abandoned croplands. Furthermore, the application of organic manures may cause 340 a short-lived increase in enzyme activity from inputs of faecal bacteria in fresh animal dung, or 341 by promoting the secretion of enzymes by Gram positive bacteria, including Actinobacteria, in 342 particular (Zhang et al., 2015b). In this study, soil oxidase and hydrolase activities were 343 observed to display coincident activity as vegetative succession progressed, associated with 344 increasing decomposition of greater organic matter inputs from larger NPP, which was 345 consistent with the results of Zhang et al., (2012). Soil hydrolases and oxidases often co-occur 346 in the same vegetation type (Baldrian, 2014), because (1) the consumption of peroxide and 347 phenolic substances by oxidases may alleviate the inhibitory effects of phenolic molecules on 348 soil enzymes and microorganisms (Sinsabaugh, 2010), which may, in turn, promote soil 349 hydrolase activities; and (2) soil oxidative enzymes may also increase substrate supply for 350 hydrolase activity by degrading complex polymeric organic matter, which could promote 351 further increases in soil hydrolase activities and nutrient contents (i.e. positive feedback) (Tian 352 & Shi, 2014). Therefore, intensified oxidase activities may indicate increased biological 353 availability of nutrients in secondary and primary forests that promotes further microbial 354 activity (Li et al., 2018). The increased SOC and DOC contents in the secondary and primary 355 forests (Table 3) may trigger soil oxidase activities as SOC and DOC supply energy to facilitate

the energy-intensive process of soil oxidase degradation (Schimel & Weintraub, 2003).
Increased inputs of complex phenolic substances, i.e. lignin and suberin, may promote increases
in oxidase activity in secondary and primary forests (Nannipieri et al., 2012).

359 Enzyme activities were reduced below 30 cm compared to the top 30 cm of the soil profile, 360 which reflects the lack of biologically available C and nutrients in deeper soils horizons 361 (Fontaine et al., 2007; Schrumpf et al., 2013) and reduced porosity and pore connectivity which 362 hinders the diffusion of water, enzymes and substrates and oxygen (Dungait et al., 2012b; Bai 363 et al., 2015). Furthermore, and consistent with previous studies (e.g. Peng & Wang, 2016; Stone 364 et al., 2014), land use (vegetation) type had a predominant effect on the soil enzyme activities 365 in the top 30 cm of the soil profile, corresponding to changes in the distribution of litter, roots 366 and SOM (Jackson et al., 1996; Lee et al., 2014). A meta-analysis of the root distribution in all 367 ecosystems averaged worldwide showed that approximately 65% of the plant roots were 368 distributed in the top 30 cm of the soil profile (Jackson et al., 1996). Roots supply nutrients 369 through root turnover, rhizoexudation and the creation of semi-permanent soil pores allowing 370 improved infiltration of water, nutrients and oxygen from the soil surface. This is disrupted in 371 intensively managed agricultural systems through tillage and harvest including removal of both 372 surface residues and roots.

373

4.2 The ratios of soil enzyme activities at different stages of vegetation restoration in the soilprofile.

376 Soil enzyme ratios provide proxy information about nutrient availability to soil microorganisms 377 and other limiting factors of the ecosystem. The C:N:P acquisition ratio 1:1.2:1.4 measured by 378 InBG:In(NAG+LAP):InAP was less than the global and regional averages (1:1:1) reported by 379 Sinsabaugh et al. (2008), which indicates that soil N and P nutrient availabilities in this 380 subtropical karst ecosystem were generally poorer than the global average. The cropland and post-abandonment land use types were all deficient in P (TP < 400 mg kg⁻¹ and available P < 4 381 mg kg⁻¹) and N (AN < 60 mg kg⁻¹) according to the classification values for Chinese soils 382 383 (National Soil Survey Office, 1998; Yang et al., 2014).

384 In agreement with our first hypothesis, and the findings of Zhang et al. (2015b), the lower ratios 385 of lnBG:ln(NAG+LAP) and the higher ratios of ln(NAG+LAP):lnAP indicated N-deficiency 386 predominated in the abandoned croplands and shrublands, which is also suggested by the 387 positive relationship indicated by lnBG:ln(NAG+LAP) and TN or AN (Table 3). Zhang et al. 388 (2015a) had also observed N-limitation during early phases of vegetation restoration in karst 389 ecosystems using changes in foliage N:P ratios. Soil N can recover relatively quickly after 390 initial rapid losses following cessation of N-fertilizer application during agricultural 391 abandonment if weathering of N-rich bedrock proceeds rapidly, or atmospheric N deposition 392 from local industrial/urban emissions are abundant, or biological N fixation (free-living or 393 rhizobial) develops (Wen et al., 2016, Li et al., 2018). However, N-loss from karstic soils 394 through runoff and leaching may be relatively rapid because of the shallow soil depth, and 395 potential for accelerated losses through the porous bedrock and the epikarst that maintains soil 396 N deficiencies in the long term (Song et al., 2017). Subsequently, we observed that the poor 397 AN did not change within the soil profiles, coincident with an increase in N-acquiring hydrolase 398 activities because N was scarce, in accordance with enzyme economic theory (Sinsabaugh et 399 al., 2002).

400 We also hypothesized that the productivity of soils in karst ecosystems would be P-limited 401 because of the paucity of the element in the parent rock and strong adsorption to Ca minerals. 402 We observed that the lnBG:lnAP and ln(NAG+LAP):lnAP ratios were less than the global 403 average reported by Sinsabaugh et al. (2008), indeed indicating P deficiency in the abandoned 404 croplands, shrublands, and secondary forests. In addition, inorganic P applications to sloping 405 croplands may inhibit soil AP activities (Zhang et al., 2015b). By contrast, AP activity in the 406 primary forests was decreased relative to C-acquiring enzyme activities in the primary forest at 407 Tianlong Mountain. Unlike P-limitation observed in the primary forest by Zhang et al. (2015a), 408 we suggest that the natural forest ecosystem studied herein was well-adapted to P-deficiency 409 through the development of mutualistic relationships between trees and arbuscular mycorrhizal 410 fungi, for example (Liang et al., 2016). The larger soil microbial biomass in undisturbed 411 primary forests in karst ecosystems (Zhu et al., 2012), relative to grasslands or shrublands, is 412 likely adapted to alleviate P limitation by releasing hydrolase enzymes to catalyze SOM 413 decomposition and promote nutrient recycling. Furthermore, rhizoexudates, including organic 414 acids, are more abundant in forests than shrublands (Jackson et al., 1996), increasing the 415 potential for the acquisition of P bound to Ca minerals and from SOM (Pan et al., 2016). the TP 416 and available P contents of the sloping and abandoned croplands, shrublands and secondary 417 forests were considered 'deficient' with respect to the classification of soil P contents across 418 China (National Soil Survey Office, 1998; Yang et al. 2014), i.e. class 5, soil TP (200-400 mg 419 kg^{-1}) and available P content (3-5 mg kg^{-1}). From the perspective of nutrient ratios, in addition, 420 the soil C:P and N:P ratios indicated that the P deficiency was more severe in the shrublands 421 and secondary forests than in the other land use types (Tian et al., 2010). In the abandoned 422 croplands, the available P concentrations were relatively poor, which might indicate that the 423 measured variations in nutrient stoichiometry were a less sensitive indicator of biological 424 nutrient availability than soil enzyme ratios.

425

426 5 CONCLUSIONS

In fragile and nutrient-limited karst ecosystems, vegetation restoration after agricultural 427 428 abandonment caused changes in soil enzyme activities in the top 0-30 cm soil depth, indicating 429 that the nutrient (C, N and P) acquisition abilities of soil microorganisms adapted to the 430 changing soil environment at each stage of recovery towards climax vegetation (forest). 431 Oxidase activities, soil nutrient and soil water contents indicated improved nutrient availability 432 and soil quality in the secondary forests, which was similar to the primary forests. The 433 InBG:In(NAG+LAP) and InBG:InAP ratios indicated that soil N was deficient in the abandoned 434 croplands and shrublands, and that soil P was deficient in the shrublands and secondary forests. 435 These results suggest that soil microorganisms in karst systems have the potential to adapt to 436 nutrient limitations and subsequently alleviate nutrient limitations to vegetation through their 437 scavenging activities, including symbiotic relationships with mycorrhiza. However, we suggest 438 that ecosystem recovery might be accelerated by the judicious application and monitoring of N 439 and P by directly alleviating nutrient limitation as part of a national strategy to rejuvenate 440 degraded and nutrient limited karst ecosystems in southwest China under GPP.

441

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450 CONFLICT OF INTEREST STATEMENT

451 The authors declare no conflicts of interest.

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Tables

| Land use type | Dominant species | Slope (°) | Altitude (m) | Disturbance history |
|--------------------|--|-----------|--------------|--|
| Sloping cropland | Zea mays | 5-10 | 1421-1435 | Maize rotated with soybeans and rape seed oil. Animal or human excreta combined with urea or compound fertilizer applied at regular intervals. |
| Abandoned cropland | Conyza canadensis, Artemisia dubia, | 12-31 | 1405-1431 | Abandoned for ~2-3 years. |
| Shrubland | Rubusparvifoliu sp., Rubusinopertus sp., Litsearubescen sp., Rosacymosa sp., Artemisia sp., Rhus chinensis | 38-49 | 1447-1484 | Recovering naturally for ~2-3 years. |
| Secondary forest | Rhus chinensis, Litsea rubescens Lecomte., Populus adenopoda Maxim., Toona sinensis (A.Juss.) Roem. | 40-46 | 1449-1471 | Recovering naturally for ~ 10 years |
| Primary forest | Itea yunnanensis, Carpinus pubescens, Lithocarpus confinis | 33-40 | 1421-1503 | More than 100 years without disturbance. |

Table 1 Characteristics of the sampled sloping cropland, abandoned cropland, shrubland, secondary forest, and primary forest.

| Soil properties | Depth (cm) | Sloping cropland | Abandoned cropland | Shrub land | Secondary forest | Primary forest |
|----------------------------|---------------|---------------------------|--------------------------|---------------------------|-------------------------|------------------------|
| | 0-10 | $36\pm1^{\rm B}$ | $41\pm2^{\text{AB}}$ | 46 ± 9^{AB} | 63 ± 3^{Aa} | $61\pm3^{\mathrm{Aa}}$ |
| | 10-30 | $34\pm2^{\text{B}}$ | $38\pm2^{\text{A}}$ | $41\pm1^{\text{AB}}$ | 50 ± 4^{Aab} | 45 ± 3^{ABb} |
| SWC (%) | 30-50 | 34 ± 2 | 39 ± 2 | 38 ± 3 | 40 ± 2^{bc} | 39 ± 2^{b} |
| | >50 | 35 ± 1 | _ | - | 34 ± 2^{c} | 38 ± 3^{b} |
| | 0-10 | 6.8 ± 0.5 | 7.2 ± 0.2 | 6.9 ± 0.1 | 7.0 ± 0.3 | 6.6 ± 0.5 |
| ŤŤ | 10-30 | 6.9 ± 0.2 | 7.3 ± 0.2 | 6.7 ± 0.2 | 7.0 ± 0.2 | 7.1 ± 0.3 |
| pH | 30-50 | 7.2 ± 0.3 | 7.4 ± 0.2 | 7.1 ± 0.1 | 7.2 ± 0.1 | 7.2 ± 0.3 |
| | >50 | 7.1 ± 0.1 | - | - | 7.7 ± 0.2 | 7.1 ± 0.3 |
| | 0-10 | $24.7\pm1.45^{\text{Da}}$ | $30.3\pm2.67^{\rm D}$ | $46.7\pm5.16^{\text{Ca}}$ | 64.7 ± 0.61^{Ba} | $77.9\pm5.61^{\rm Aa}$ |
| | 10-30 | 21.0 ± 0.84^{Ba} | $27.6\pm4.27^{\text{B}}$ | 35.7 ± 1.84^{ABa} | $51.3\pm6.43^{\rm Aa}$ | $47.6\pm9.40^{\rm Ab}$ |
| SOC (g kg ⁻¹) | 30-50 | $11.9\pm2.98^{\text{Bb}}$ | $24.1\pm4.93^{\rm AB}$ | 20.8 ± 2.37^{ABb} | 27.6 ± 4.53^{ABb} | 35.7 ± 2.97^{Abc} |
| | >50 | $8.0\pm1.64^{\text{Bb}}$ | - | - | 15.4 ± 2.21^{ABb} | $21.8\pm3.35^{\rm Ac}$ |
| | 0-10 | $29\pm2.2^{\text{Ba}}$ | $49\pm9.0^{\rm B}$ | $138\pm16.1^{\rm A}$ | $85\pm21.6^{\text{AB}}$ | $140\pm27.4^{\rm Aa}$ |
| DOC (mg kg ⁻¹) | 10-30 | 24 ± 2.0^{ab} | 39 ± 7.6 | 80 ± 24.3 | 93 ± 26.3 | 73 ± 7.6^{ab} |

Table 2 Mean ($n=4, \pm 1$ standard deviation) soil properties at each depth for each land use. Definition of abbreviations: SWC: soil water content, SOC: soil organic carbon, DOC: dissolved organic carbon, TN: total nitrogen, AN: available nitrogen, TP: total phosphorus, available P: available phosphorus.

| | 30-50 | 22 ± 3.1^{ab} | 28 ± 2.8 | 59 ± 22.8 | 44 ± 8.3 | 68 ± 17.5^{ab} |
|----------------------------|-------|--------------------------|--------------------------|--------------------------|--------------------------|---------------------------|
| | >50 | $17\pm1.5^{\mathrm{b}}$ | - | _ | 34 ±11.1 | 40±6.1 ^b |
| | 0-10 | $2.5\pm0.22^{\text{Ba}}$ | $3.1\pm0.18^{\text{B}}$ | $4.5\pm0.47^{\text{Ba}}$ | $4.5\pm1.04^{\text{Ba}}$ | $7.3\pm0.33^{\text{Aa}}$ |
| $TN(\alpha k \alpha^{-1})$ | 10-30 | 2.1 ± 0.09^{Bab} | $2.7\pm0.30^{\text{B}}$ | 3.1 ± 0.08^{ABb} | $4.4\pm0.49^{\text{Aa}}$ | $4.6\pm0.48^{\text{Aab}}$ |
| TN (g kg ⁻¹) | 30-50 | $1.5\pm0.23^{\text{Bb}}$ | $2.4\pm0.44^{\text{AB}}$ | $2.1\pm0.13^{\text{Bb}}$ | 2.7 ± 0.41^{ABab} | 3.7 ± 0.15^{Ab} |
| | >50 | $1.5\pm0.39^{\text{b}}$ | - | - | $1.8\pm0.21^{\text{b}}$ | 3.7 ± 1.11^{b} |
| | 0-10 | $40\pm6.1^{\text{BC}}$ | $24\pm3.7^{\rm C}$ | $48\pm4.8^{\rm AB}$ | 57 ± 7.0^{ABa} | $70\pm3.0^{\text{Aa}}$ |
| | 10-30 | 30 ± 5.1 | 23 ± 2.4 | 52 ± 13.1 | 37 ± 7.8^{ab} | $31 \pm 1.9^{\text{b}}$ |
| AN (mg kg ⁻¹) | 30-50 | $25\pm 6.0^{\text{B}}$ | $27\pm3.8^{\text{B}}$ | $49\pm4.8^{\rm A}$ | $25\pm1.8^{\text{Bb}}$ | $23\pm1.8^{\text{Bbc}}$ |
| | >50 | 22 ± 1.3 | - | - | $21\pm5.1^{\text{b}}$ | $20\pm2.5^{\rm c}$ |
| | 0-10 | $487\pm29^{\text{Ba}}$ | $503\pm42^{\text{Ba}}$ | $348\pm22^{\text{Ba}}$ | $377\pm26^{\text{Ba}}$ | $1180 \pm 173^{\rm A}$ |
| | 10-30 | 339 ± 39^{Bab} | 404 ± 36^{Bab} | $275\pm4^{\text{Bb}}$ | 323 ± 32^{Bab} | $841\pm 66^{\rm A}$ |
| $TP (mg kg^{-1})$ | 30-50 | $271\pm46^{\text{Bb}}$ | $332\pm38^{\text{Bb}}$ | $229 \pm 11^{\text{Bb}}$ | 222 ± 25^{Bb} | $765\pm93^{\rm A}$ |
| | >50 | $276\pm 46^{\text{Bb}}$ | - | - | $221\pm14^{\text{Bb}}$ | $643 \pm 132^{\rm A}$ |
| | 0-10 | $4\pm0.8^{\rm A}$ | $2\pm0.4^{\rm B}$ | $4\pm0.8^{\rm AB}$ | $4\pm0.8^{\rm AB}$ | $9\pm2.8^{\rm A}$ |
| | 10-30 | $3\pm1.2^{\rm B}$ | $1\pm0.3^{\text{B}}$ | $4\pm0.8^{\rm B}$ | $2\pm0.8^{\text{B}}$ | $12\pm3.5^{\rm A}$ |
| AP (mg kg ⁻¹) | 30-50 | $3\pm0.7^{\rm A}$ | $1\pm0.7^{\mathrm{B}}$ | $3\pm0.5^{\rm AB}$ | $2\pm0.6^{\text{AB}}$ | $7\pm2.4^{\rm A}$ |
| | >50 | $3\pm0.8^{\text{B}}$ | - | - | $3\pm0.5^{\rm B}$ | $12\pm3.7^{\rm A}$ |

| | 0-10 | $50.8\pm2.35^{\text{Bab}}$ | $61.8\pm8.24^{\rm B}$ | $137.6\pm19.84^{\mathrm{Aa}}$ | $173.8\pm11.22^{\text{Aa}}$ | $72.0\pm15.15^{\text{Ba}}$ |
|--------|-------|----------------------------|-------------------------|-------------------------------|-----------------------------|----------------------------|
| SOCIED | 10-30 | $64.3\pm7.31^{\rm Ba}$ | $68.5\pm9.40^{\rm B}$ | $130.3\pm9.40^{\text{Aab}}$ | $165.7\pm29.95^{\text{Aa}}$ | 56.7 ± 10.08^{Bab} |
| SOC:TP | 30-50 | 42.6 ± 4.81^{Bbc} | $70.3\pm10.54^{\rm B}$ | $90.4\pm7.59^{\text{ABb}}$ | 124.6 ± 18.78^{Aab} | $49.1\pm7.66^{\text{Bab}}$ |
| | >50 | $29.5\pm4.80^{\text{Bc}}$ | _ | - | 71.1±12.90 ^{Ab} | $35.3\pm3.47^{\text{Bb}}$ |
| | 0-10 | $5.2\pm0.45^{\text{B}}$ | $6.2\pm0.64^{\text{B}}$ | $12.9\pm0.57^{\rm A}$ | 11.8±2.76 ^{AB} | $6.7\pm1.24^{\text{AB}}$ |
| | 10-30 | 6.6 ± 0.80^{BC} | 6.6 ± 0.63^{BC} | $11.4\pm0.33^{\rm AB}$ | 14.2±2.36 ^A | $5.5\pm0.46^{\rm C}$ |
| TN:TP | 30-50 | $5.6\pm0.49^{\rm B}$ | $7.2\pm0.82^{\rm AB}$ | $9.0\pm0.25^{\text{AB}}$ | $12.6\pm2.37^{\rm A}$ | $5.0\pm0.39^{\text{B}}$ |
| | >50 | 5.3 ± 0.51 | - | - | 8.1 ± 1.26 | 5.6 ± 1.28 |

Superscripted uppercase letters indicate significant different sub-groups between land uses at p < 0.05 in each soil layer, while lowercase letters indicate significant different depth sub-groups in each land use.

| | SWC | pН | SOC | DOC | TN | AN | TP | available P | SOC:TN | SOC:TP | TN:TP |
|------------------|---------|---------|---------|---------|---------|---------|---------|-------------|---------|--------|--------|
| BG | 0.705** | -0.272* | 0.730** | 0.526** | 0.653** | 0.549** | 0.448** | 0.372** | 0.355** | 0.275* | 0.122 |
| NAG | 0.639** | -0.171 | 0.723** | 0.555** | 0.632** | 0.592** | 0.569** | 0.391** | 0.340** | 0.197 | 0.030 |
| BX | 0.673** | -0.154 | 0.688** | 0.456** | 0.638** | 0.525** | 0.509** | 0.318** | 0.320** | 0.181 | 0.035 |
| CBH | 0.551** | -0.217 | 0.585** | 0.412** | 0.522** | 0.505** | 0.462** | 0.238* | 0.268* | 0.099 | -0.046 |
| LAP | 0.501** | -0.122 | 0.462** | 0.411** | 0.389** | 0.306** | 0.373** | 0.221 | 0.241* | 0.095 | -0.050 |
| AP | 0.638** | -0.234* | 0.635** | 0.570** | 0.556** | 0.549** | 0.444** | 0.317** | 0.300* | 0.250* | 0.118 |
| PPO | 0.486** | -0.103 | 0.564** | 0.282* | 0.544** | 0.326** | 0.373** | 0.353** | 0.257* | 0.219 | 0.129 |
| POD | 0.470** | -0.224 | 0.543** | 0.338** | 0.579** | 0.348** | 0.515** | 0.204 | 0.159 | 0.082 | 0.009 |
| InBG:ln(NAG+LAP) | 0.321** | -0.289* | 0.388** | 0.276* | 0.358** | 0.252* | 0.239* | 0.160 | 0.195 | 0.135 | 0.194 |
| ln(NAG+LAP):InAP | 0.195 | 0.202 | 0.189 | 0.072 | 0.146 | 0.015 | 0.203 | -0.097 | 0.083 | -0.108 | -0.014 |
| InBG:InAP | 0.479** | -0.122 | 0.553** | 0.339** | 0.490** | 0.282* | 0.412** | 0.096 | 0.268* | 0.06 | 0.194 |

Table 3 Pearson correlation analyses of the soil enzyme activities and soil properties.

* and ** indicate the significant difference at p < 0.05 and p < 0.01. Definition of abbreviations: SWC: soil water content, SOC: soil organic carbon, DOC: dissolved organic carbon, TN: total nitrogen, AN: available nitrogen, TP: total phosphorus, available P: available phosphorus, BG: β -1,4glucosidase, CBH: β -D-1,4-cellobiosidase, BX: β -1,4-xylosidase, NAG: β -1,4-N-acetyl-glucosaminidase, LAP: Leucine aminopeptidase, AP: alkaline phosphatase, PPO: Polyphenol oxidase, POD: Peroxidase.

Figure captions

Fig. 1 Soil enzyme activities in soil profile under different land uses. Uppercase letters above error bars indicate significant differences between land use sub-sets at p < 0.05 in each soil depth, and lowercase letters indicate significant different soil depth sub-sets in each land use. Additional notation as follows - BG: β -1,4-glucosidase, CBH: β -D-1,4-cellobiosidase, BX: β -1,4-xylosidase, NAG: β -1,4-N-acetyl-glucosaminidase, LAP: Leucine aminopeptidase, AP: alkaline phosphatase, PPO: Polyphenol oxidase, POD: Peroxidase.

Fig. 2 Comparison of soil enzyme activity ratios among the land use types. The lnBG:ln(NAG+LAP) indicates the ratios of natural logarithm of C (BG) to natural logarithm of N (NAG+LAP) nutrient acquiring hydrolase activities; The lnBG:lnAP indicates the ratios of natural logarithm of C (BG) to natural logarithm of P (AP) nutrient acquiring hydrolase activities; ln(NAG+LAP):lnAP indicates the ratios of natural logarithm of N (NAG+LAP) to natural logarithm of P (AP) nutrient acquiring hydrolase activities. Uppercase letters indicate significant different between land uses at p < 0.05.

Fig. 3 Comparison of soil enzyme activity ratios among the depths. Lowercase letters indicate significant different between soil depths at p < 0.05

Supplementary Tables

Table S1 Functions and substrates of the enzymes quantified in this study. EC: enzyme commission, MUB: 4-Methylumbelliferone.

Table S2 Two-way ANOVA analysis of the soil properties. *, ** and *** indicate the significant difference at p < 0.05, p < 0.01 and p < 0.001.

Table S3 Comparison of soil properties among the land use types. Lowercase letters indicate significant different between land use types at p < 0.05.

Table S4 Comparison of soil properties among the soil depths.Lowercase letters indicatesignificant different between soil depths at p < 0.05.

| | Depth | Land use type | Depth ×Land use type |
|------------------|----------|---------------|----------------------|
| SWC | 15.67*** | 9.64*** | 3.08** |
| pH | 2.49 | 1.40 | 0.37 |
| SOC | 55.32*** | 33.50*** | 4.46*** |
| DOC | 10.76*** | 12.47*** | 1.84 |
| TN | 19.41*** | 21.74*** | 2.09* |
| AN | 17.22*** | 9.45*** | 3.79** |
| TP | 13.14*** | 54.36*** | 1.30 |
| available P | 0.75 | 15.31*** | 0.60 |
| SOC:TN | 3.66* | 2.62* | 0.89 |
| SOC:TP | 12.16*** | 35.42*** | 1.95 |
| TN: TP | 2.03 | 20.23*** | 1.34 |
| BG | 41.57*** | 11.33*** | 2.67* |
| BX | 20.30*** | 9.91*** | 3.08** |
| CBH | 25.90*** | 5.85*** | 2.25* |
| NAG | 48.14*** | 25.16*** | 8.69*** |
| LAP | 6.17** | 5.22** | 2.2* |
| AP | 17.56*** | 10.75*** | 2.53* |
| PPO | 43.04*** | 18.04*** | 2.39* |
| POD | 38.20*** | 12.82*** | 2.84* |
| lnBG:ln(NAG+LAP) | 15.36*** | 3.63* | 1.64 |
| ln(NAG+LAP):lnAP | 1.50 | 3.31* | 0.67 |
| lnBG:lnAP | 40.19*** | 3.96** | 1.36 |

Table S2 Two-way ANOVA of soil properties (variables) with soil depth (Depth) and Land Use Type as factors.

*, ** and *** indicate the significant difference at p < 0.05, p < 0.01 and p < 0.001. Definition of abbreviations: SWC: soil water content, SOC: soil organic carbon, DOC: dissolved organic carbon, TN: total nitrogen, AN: available nitrogen, TP: total phosphorus, available P: available phosphorus, BG: β -1,4-glucosidase, CBH: β -D-1,4-cellobiosidase, BX: β -1,4-xylosidase, NAG: β -1,4-N-acetyl-glucosaminidase, LAP: Leucine aminopeptidase, AP: alkaline phosphatase, PPO: Polyphenol oxidase, POD: Peroxidase.

| Soil properties | Slopping | Abandoned | Shrubland | Secondary | Primary |
|---|-------------------------|------------------------------|-----------------------------|------------------------------|------------------------|
| Soil properties | cropland | cropland | Sillubialiu | forest | forest |
| SWC (%) | 35±1 ^B | 39 ± 1^{AB} | 42 ± 3^{AB} | 48 ± 3^{A} | 46±3 ^A |
| SOC $(g kg^{-1})$ | $16.4 \pm 1.92^{\circ}$ | $27.3{\pm}2.26^{BC}$ | $34.4{\pm}3.67^{\text{AB}}$ | 41.4 ± 5.4^{A} | 45.8 ± 5.96^{A} |
| DOC (mg kg ⁻¹) | $23 \pm 1.5^{\circ}$ | $39 \pm 4.5^{\circ}$ | $92{\pm}15.0^{A}$ | 66±10.9 ^B | $80{\pm}12.1^{AB}$ |
| TN (g kg ⁻¹) | $1.9 \pm 0.16^{\circ}$ | $2.7{\pm}0.19^{\rm BC}$ | 3.2 ± 0.34^{B} | 3.4 ± 0.42^{B} | 4.8 ± 0.47^{A} |
| AN (mg kg ⁻¹) | 29 ± 2.9^{B} | 24 ± 1.8^{B} | 50 ± 4.5^{A} | 36 ± 4.6^{B} | 36 ± 5.3^{B} |
| TP (mg kg ⁻¹) | 343 ± 29^{BC} | 413 ± 29^{B} | 283 ± 16^{C} | 291±21 ^C | 858 ± 75^{A} |
| available P (mg kg ⁻¹) | 3 ± 0.4^{B} | 2±0.3 ^B | 3 ± 0.4^{B} | 3±0.4 ^B | $10{\pm}1.5^{A}$ |
| SOC:TN | 8.2 ± 0.53^{B} | $9.9{\pm}0.26^{\rm AB}$ | 10.8 ± 0.64^{AB} | 13.1 ± 2.30^{A} | 9.3 ± 0.55^{B} |
| SOC:TP | 46.8±3.99 ^C | $66.9{\pm}5.05^{\mathrm{B}}$ | 119.4 ± 9.27^{A} | $138.0{\pm}13.77^{A}$ | $53.3{\pm}5.64^{BC}$ |
| TN:TP | 5.7 ± 0.29^{B} | 6.7 ± 0.39^{B} | 11.1 ± 0.54^{A} | 11.9 ± 1.20^{A} | $5.7{\pm}0.45^{\rm B}$ |
| BG (nmol $h^{-1} g^{-1}$) | 148 ± 35^{B} | 110 ± 25^{B} | 99 ± 18^{B} | 271 ± 63^{AB} | 393 ± 95^{A} |
| NAG (nmol h ⁻¹ g ⁻¹) | 76 ± 22^{B} | 32 ± 7^{B} | 50 ± 8^{B} | $201{\pm}54^{AB}$ | 303 ± 78^{A} |
| BX (nmol $h^{-1} g^{-1}$) | 24 ± 5^{B} | 34 ± 6^{B} | 32 ± 4^{B} | 50 ± 9^{AB} | $80{\pm}18^{A}$ |
| CBH (nmol h ⁻¹ g ⁻¹) | 30 ± 9^{AB} | 23 ± 6^{AB} | 12 ± 2^{B} | 32 ± 8^{AB} | 58 ± 16^{A} |
| LAP (nmol $h^{-1} g^{-1}$) | 188 ± 40^{B} | 342 ± 58^{AB} | 297 ± 57^{AB} | 327 ± 54^{AB} | 484 ± 84^{A} |
| AP (nmol $h^{-1} g^{-1}$) | 898±135 ^B | 474 ± 50^{B} | 816 ± 62^{B} | 1517 ± 239^{A} | 1672±311 ^A |
| PPO (µmol h ⁻¹ g ⁻¹) | 26 ± 4^{AB} | 24 ± 3^{AB} | 20 ± 2^{B} | 36±4 ^A | 34±4 ^A |
| POD (µmol h ⁻¹ g ⁻¹) | 44 ± 4^{AB} | 32 ± 4^{B} | 32 ± 2^{B} | 46 ± 4^{AB} | 55 ± 5^{A} |
| lnBG:ln(NAG+LAP) | $0.84{\pm}0.05^{AB}$ | 0.76 ± 0.03^{B} | 0.78 ± 0.03^{B} | 0.86 ± 0.02 ^A | $0.83{\pm}0.02^{AB}$ |
| ln(NAG+LAP):lnAP | 0.81 ± 0.04^{B} | 0.95 ± 0.02^{A} | $0.84{\pm}0.03^{B}$ | $0.84{\pm}0.02^{B}$ | $0.88{\pm}0.02^{AB}$ |
| lnBG:lnAP | 0.67 ± 0.03^{B} | $0.72{\pm}0.03^{AB}$ | $0.65{\pm}0.03^{\rm B}$ | $0.72{\pm}0.02^{AB}$ | 0.73 ± 0.03^{A} |

 Table S3 Comparison of soil properties among the land use types.

Superscripted uppercase letters indicate significant different between land use types at p < 0.05.

| Table 54 Comparison of a | son properties e | unong the son t | iepuis. | |
|---|---------------------|-----------------------|-------------------------|------------------------|
| Soil properties | 0-10 cm | 10-30 cm | 30-50cm | > 50 cm |
| SWC (%) | 49±3 ^a | 42 ± 2^{b} | 38±1 ^b | 36±1 ^b |
| SOC (g kg ⁻¹) | 48.9 ± 4.85^{a} | 36.6 ± 3.43^{a} | 24±2.32 ^b | 15.1±2.29 ^b |
| DOC (mg kg ⁻¹) | 88 ± 12.5^{a} | 62 ± 8.9^{b} | 44 ± 6.7^{bc} | $30 \pm 4.6^{\circ}$ |
| TN (g kg ⁻¹) | 4.4 ± 0.44^{a} | 3.4 ± 0.26^{ab} | 2.5 ± 0.21^{b} | $2.4{\pm}0.5^{b}$ |
| AN (mg kg ⁻¹) | 48±4.1 ^a | 34±3.7 ^b | 30 ± 2.7^{bc} | 21±1.5° |
| TP (mg kg ⁻¹) | 579 ± 78^{a} | 436 ± 50^{b} | 364±51 ^b | 395 ± 76^{b} |
| SOC:TN | 12.2 ± 1.72^{a} | 10.6 ± 0.27^{a} | $9.5{\pm}0.45^{ab}$ | 7.1 ± 0.86^{b} |
| SOC:TP | 8.6 ± 0.92^{a} | 8.9 ± 0.90^{a} | $7.9{\pm}0.78^{ab}$ | 6.2 ± 0.66^{b} |
| BG (nmol $h^{-1} g^{-1}$) | 451±69 ^a | 226 ± 40^{b} | 55±6° | 53±9° |
| NAG (nmol h ⁻¹ g ⁻¹) | 312±64 ^a | 152 ± 33^{b} | 24±4° | 27±7° |
| BX (nmol $h^{-1} g^{-1}$) | 82±15 ^a | 39 ± 5^{b} | 23±3 ^b | 30±7 ^b |
| CBH (nmol h ⁻¹ g ⁻¹) | 71±13 ^a | 22 ± 4^{b} | 13±2 ^b | 16±3 ^b |
| LAP (nmol h ⁻¹ g ⁻¹) | 468 ± 70^{a} | 349 ± 37^{ab} | 211±44 ^b | 250 ± 63^{b} |
| AP (nmol $h^{-1} g^{-1}$) | 1698 ± 254^{a} | 1227 ± 200^{ab} | 611±43 ^b | 788 ± 84^{b} |
| PPO (µmol h ⁻¹ g ⁻¹) | 40±3 ^a | 32±3 ^a | 18±2 ^b | 21±3 ^b |
| POD (μ mol h ⁻¹ g ⁻¹) | 58 ± 4^{a} | 46±3 ^b | 26±2° | 40±3 ^b |
| lnBG:ln(NAG+LAP) | 0.93±0.02ª | $0.85 {\pm} 0.03^{b}$ | 0.74±0.03° | 0.72±0.05° |
| lnBG:lnAP | 0.83 ± 0.02^{a} | 0.74 ± 0.02^{b} | $0.60 \pm 0.02^{\circ}$ | 0.59±0.03° |
| | | | | |

Table S4 Comparison of soil properties among the soil depths.

Superscripted lowercase letters indicate significant different between soil depths at p < 0.05.